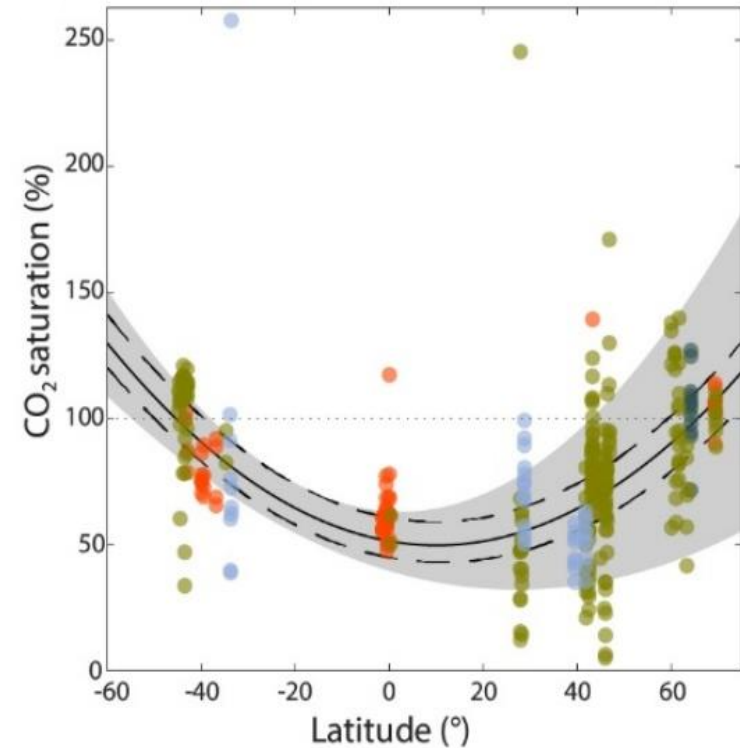


**Most streams are CO<sub>2</sub> emitters, but not all!**

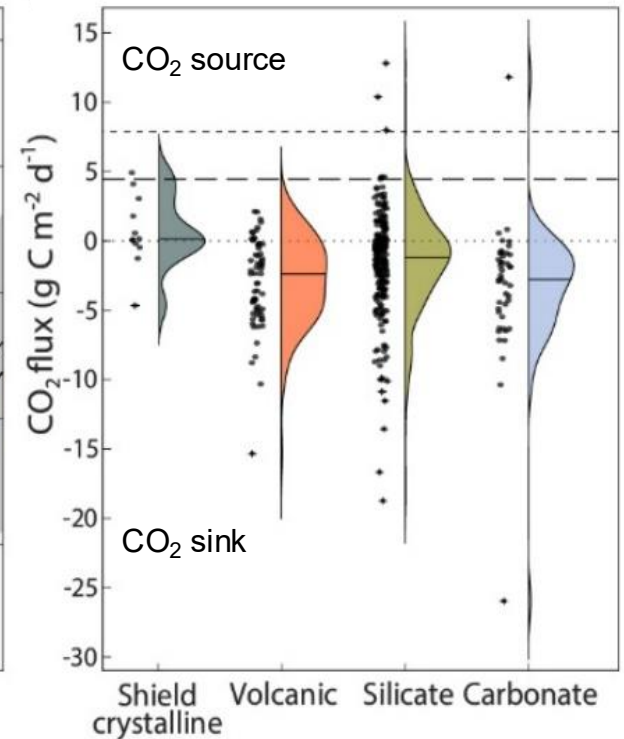
# Glacier-fed streams and lakes

- Glacier-fed streams are CO<sub>2</sub> sinks depending on parent geology
- Chemical weathering consumes CO<sub>2</sub>
- Likely in equilibrium with the atmosphere before industrialisation



● Shield crystalline  
● Volcanic  
● Silicate  
● Carbonate

— Quadratic model fit  
— 95% C.I.  
Quartiles of model fit



..... High-latitude streams areal  
CO<sub>2</sub> flux  
- - - Temperate streams areal CO<sub>2</sub> flux

# Streams and rivers are major components of the global

Review

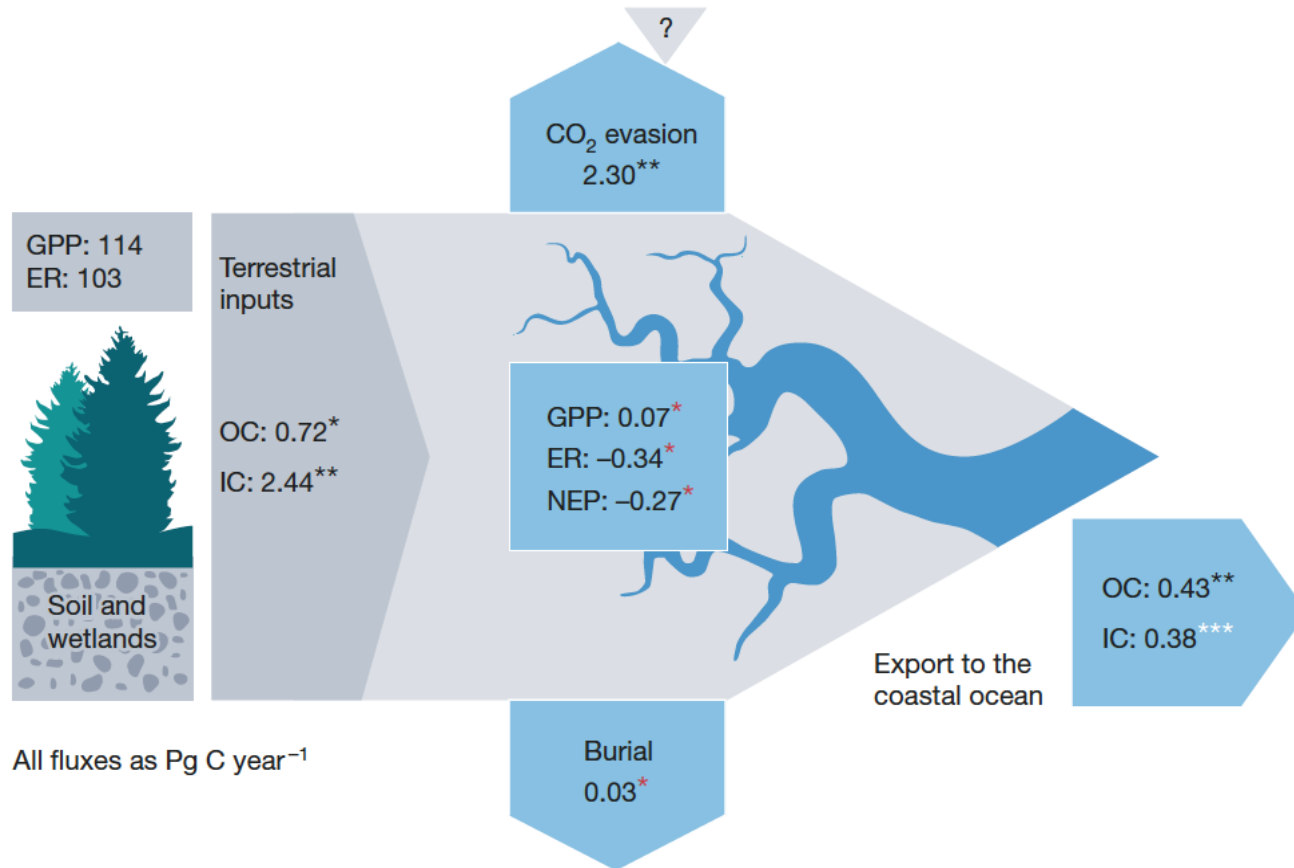
## River ecosystem metabolism and carbon biogeochemistry in a changing world

<https://doi.org/10.1038/s41586-022-05500-8>

Received: 12 March 2021

Accepted: 31 October 2022

Tom J. Battin<sup>1,2</sup>, Ronny Lauerwald<sup>2</sup>, Emily S. Bernhardt<sup>3</sup>, Enrico Bertuzzo<sup>4</sup>, Lluís Gómez Gener<sup>5</sup>, Robert O. Hall Jr<sup>6</sup>, Erin R. Hotchkiss<sup>7</sup>, Taylor Maavara<sup>8</sup>, Tamlin M. Pavelsky<sup>9</sup>, Lishan Ran<sup>10,11</sup>, Peter Raymond<sup>12</sup>, Judith A. Rosentreter<sup>12,13</sup> & Pierre Regnier<sup>14</sup>









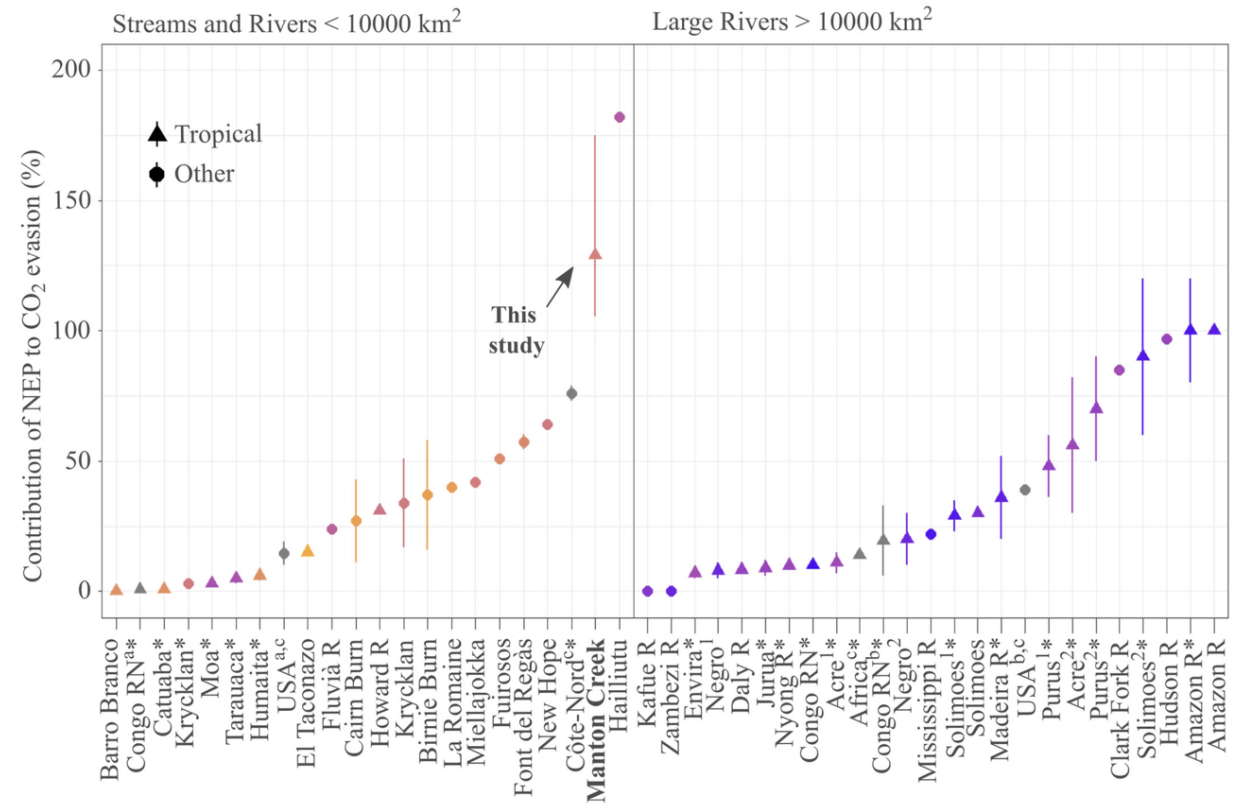
- CO<sub>2</sub> emissions from streams and rivers similar to the drawdown fluxes of atmospheric CO<sub>2</sub> by the world's oceans
- Ecosystem metabolism contributes ca. 10% to global CO<sub>2</sub> emissions – up to 90% locally.
- Remaining CO<sub>2</sub> from terrestrial primary production and chemical weathering

# River networks matter for the global carbon cycle

- Contributions from river NEP to CO<sub>2</sub> evasion from rivers vary widely
- Depend on gas exchange rate, carbonate dissolution, photochemistry etc

## Stream respiration exceeds CO<sub>2</sub> evasion in a low-energy, oligotrophic tropical stream

Vanessa Solano <sup>1,\*</sup> Clément Duvert <sup>1,2\*</sup> Christian Birkel <sup>3,4</sup> Damien T. Maher <sup>5</sup> Erica A. García <sup>1</sup>  
Lindsay B. Hutley <sup>1</sup>

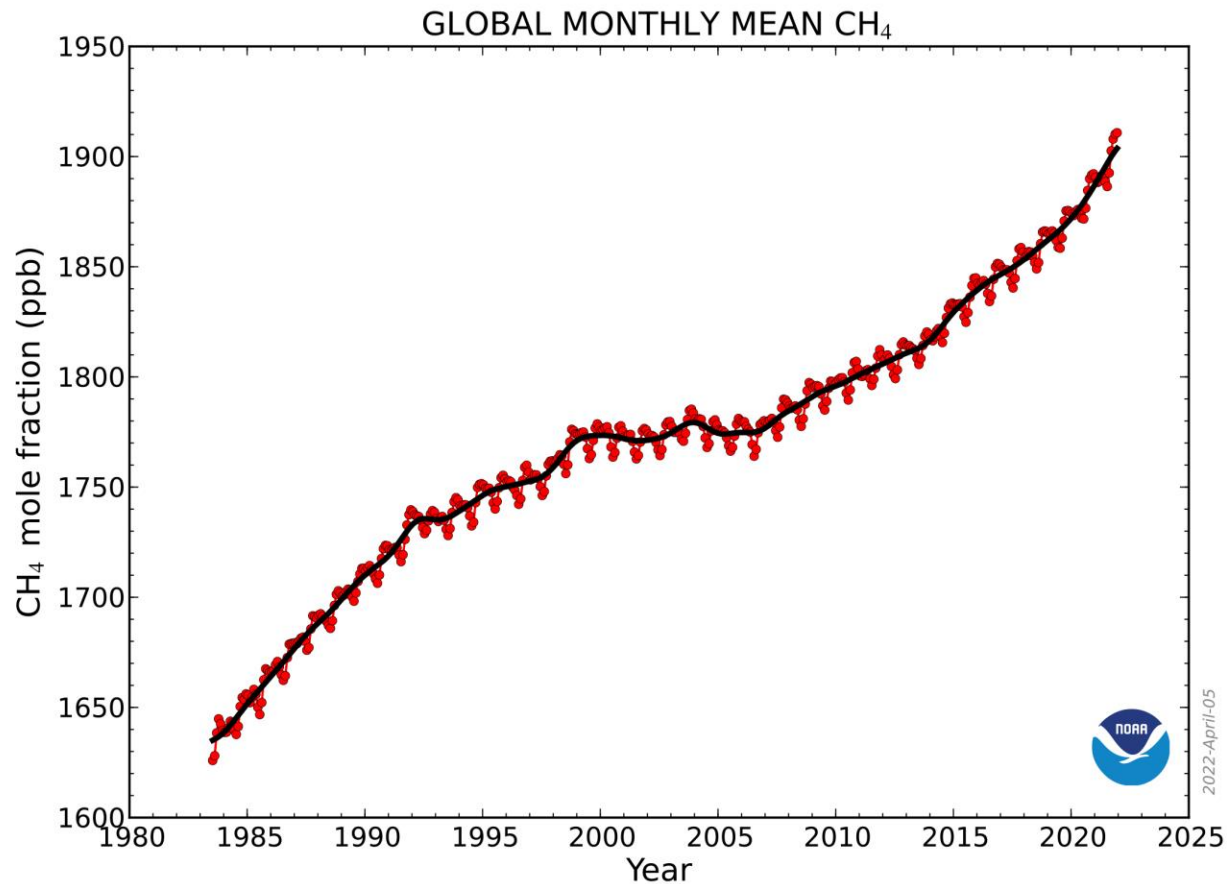


# **What about methane?**

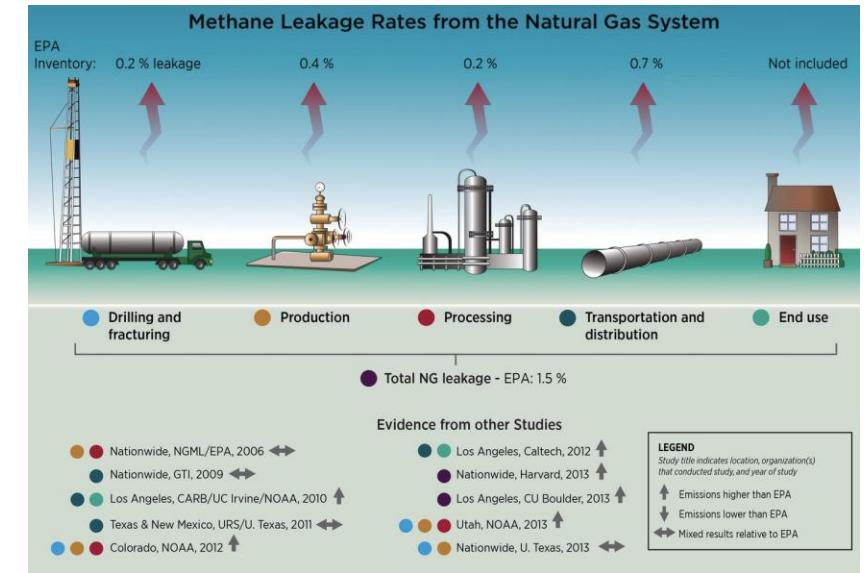
Biological sources (methanogenesis by archaea and bacteria)

Geogenic sources

Anthropogenic sources



- Methane is a potent greenhouse gas
- Its emissions are globally increasing at a rapid pace



Leaky gas infrastructures as one major source of methane to the atmosphere

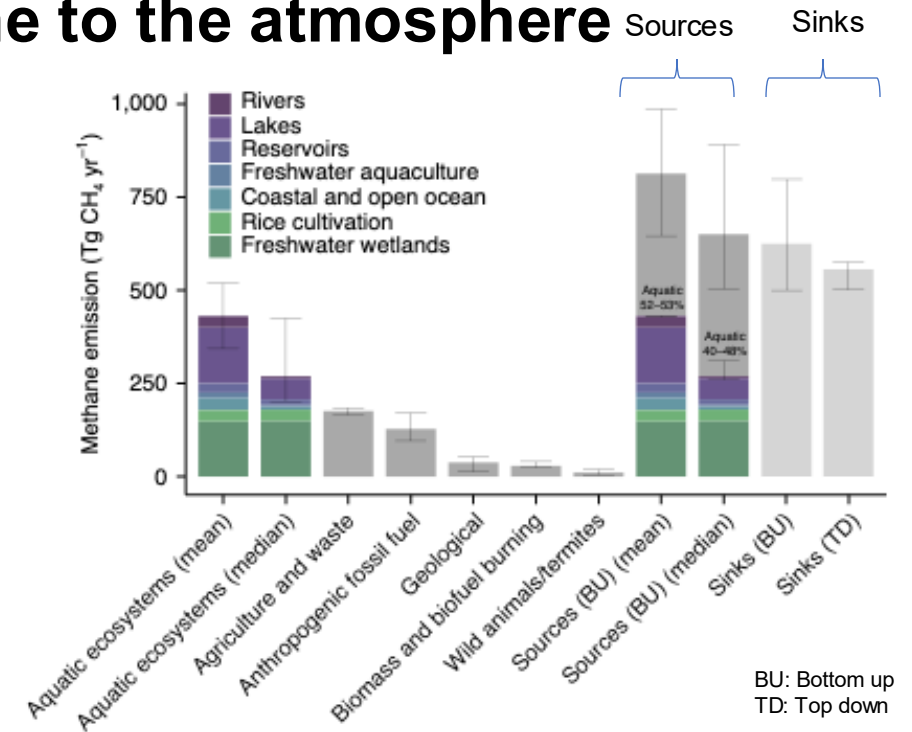
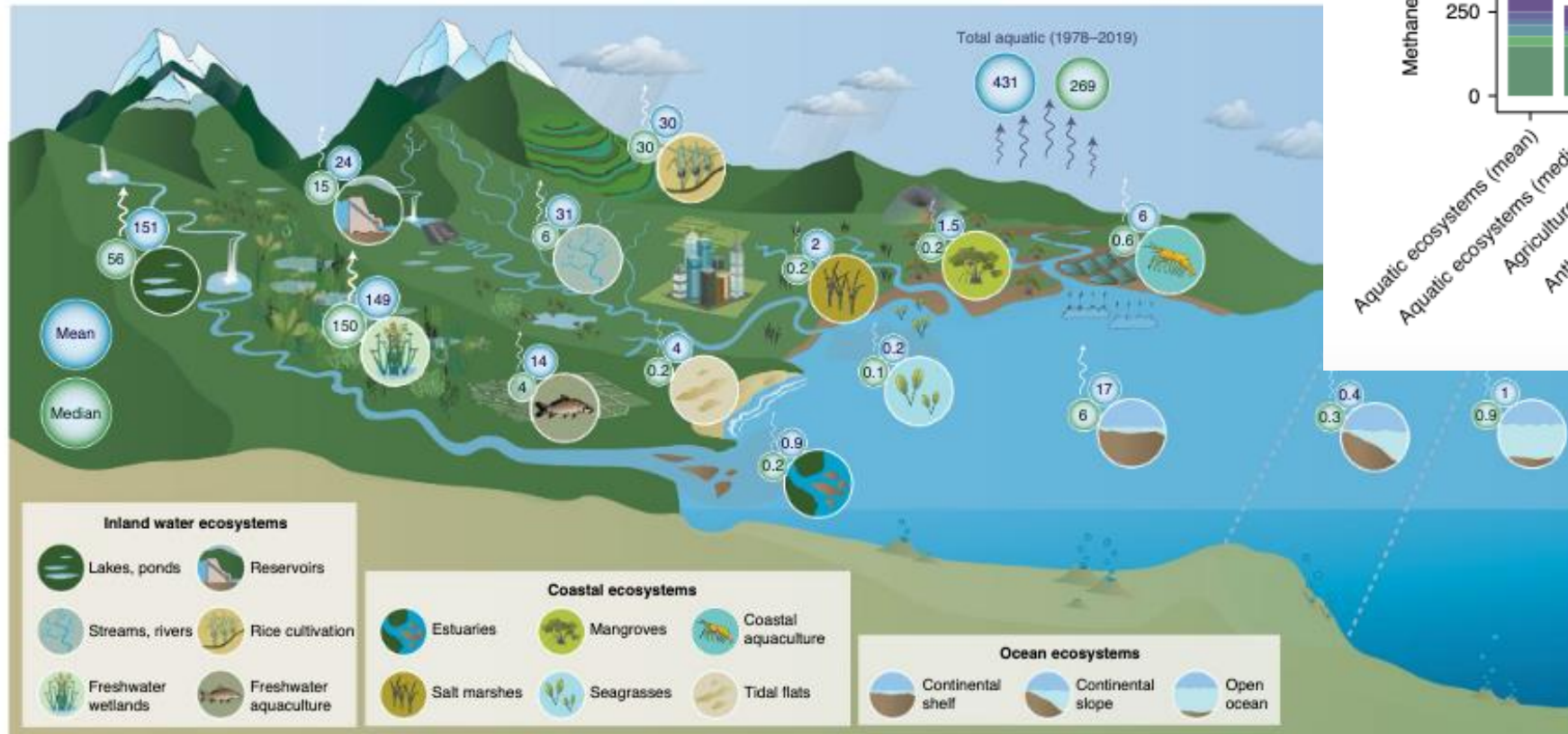
...what else?

# Streams and rivers are major sources of methane to the atmosphere



## Half of global methane emissions come from highly variable aquatic ecosystem sources

Judith A. Rosentreter<sup>1,2</sup>, Alberto V. Borges<sup>3</sup>, Bridget R. Deemer<sup>4</sup>, Meredith A. Holgerson<sup>5,6,7</sup>, Shaoda Liu<sup>2,8</sup>, Chunlin Song<sup>9,10</sup>, John Melack<sup>11</sup>, Peter A. Raymond<sup>2</sup>, Carlos M. Duarte<sup>12,13</sup>, George H. Allen<sup>14</sup>, David Olefeldt<sup>15</sup>, Benjamin Poulter<sup>16</sup>, Tom I. Battin<sup>17</sup> and Bradley D. Eyre<sup>1</sup>



$$1\text{Tg} = 10^{12} \text{ g}$$

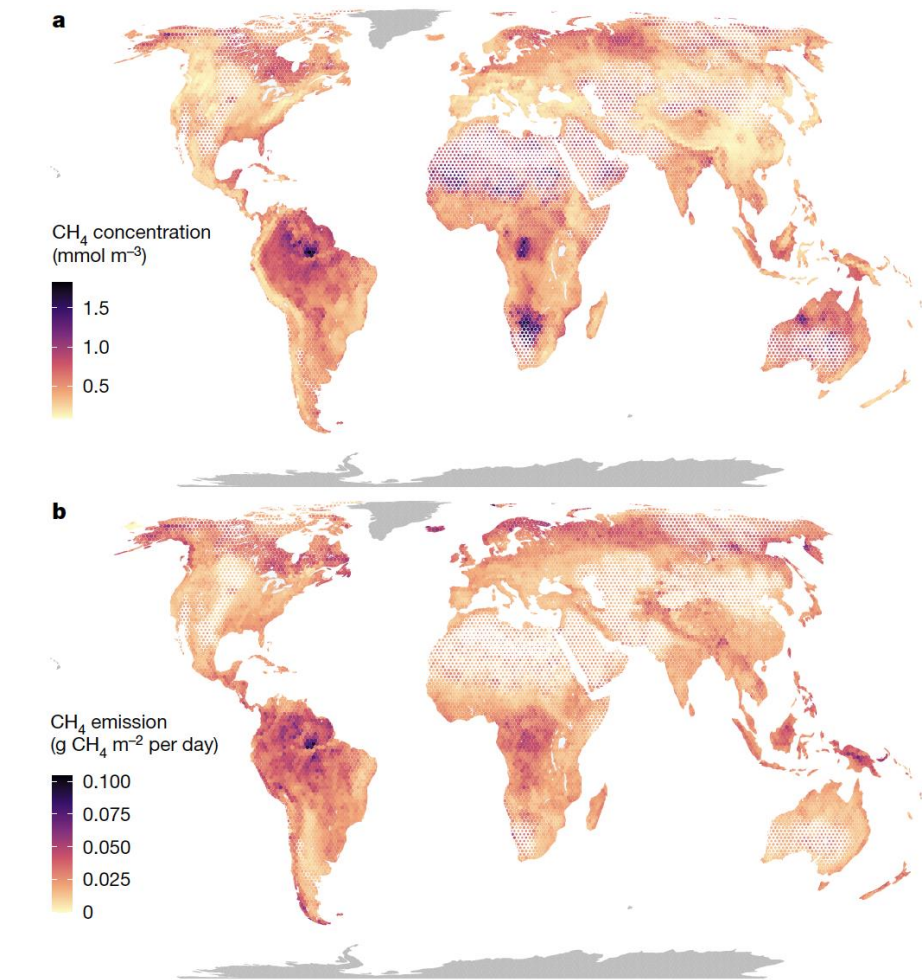
**Fig. 3 | Global aquatic methane emissions from headwater streams to the open ocean.** Numbers are Tg CH<sub>4</sub> yr<sup>-1</sup>. Mean emissions are shown in blue circles, and median emissions are shown in green circles. The relative importance of the factors controlling methane distribution and emissions varies along the land–ocean aquatic continuum.



# Global methane emissions from rivers and streams

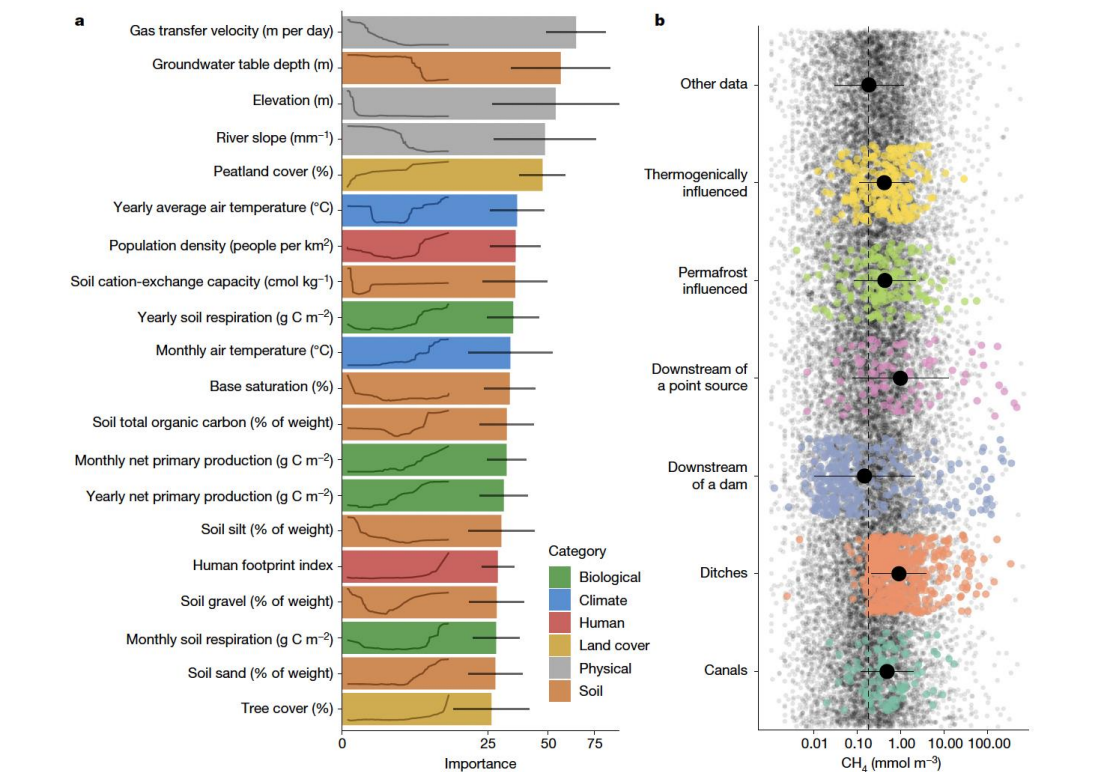
<https://doi.org/10.1038/s41586-023-06344-6>  
Received: 25 October 2022

Gerard Rocher-Ros<sup>1,2,3,✉</sup>, Emily H. Stanley<sup>4</sup>, Luke C. Loken<sup>5</sup>, Nora J. Casson<sup>6</sup>, Peter A. Raymond<sup>7</sup>, Shaoda Liu<sup>7,8</sup>, Giuseppe Amatulli<sup>7</sup> & Ryan A. Sponseller<sup>1</sup>



**Fig. 1 | Global patterns of CH<sub>4</sub> in rivers and streams. a, b.** Modelled yearly average CH<sub>4</sub> concentrations (a) and emissions (b) in rivers and streams. Data have been aggregated in hexagonal bins, and the size of each hexagon is rescaled with runoff, to better visualize patterns in areas with high coverage of running waters. Areas with runoff greater than 1,500 mm per year have full-sized hexagons; hexagons in areas with runoff of 500 mm per year have been reduced by 10%; and hexagons with a runoff less than 50 mm per year have been reduced by 50%. The model could not be applied in Greenland and Antarctica, which are shown in dark grey.

- Streams and rivers emit 27.9 (16.7–39.7) Tg CH<sub>4</sub> per year, roughly equal emissions from lakes and ponds
- Physical ecosystem attributes and land use as some of the best predictors for CH<sub>4</sub> concentrations



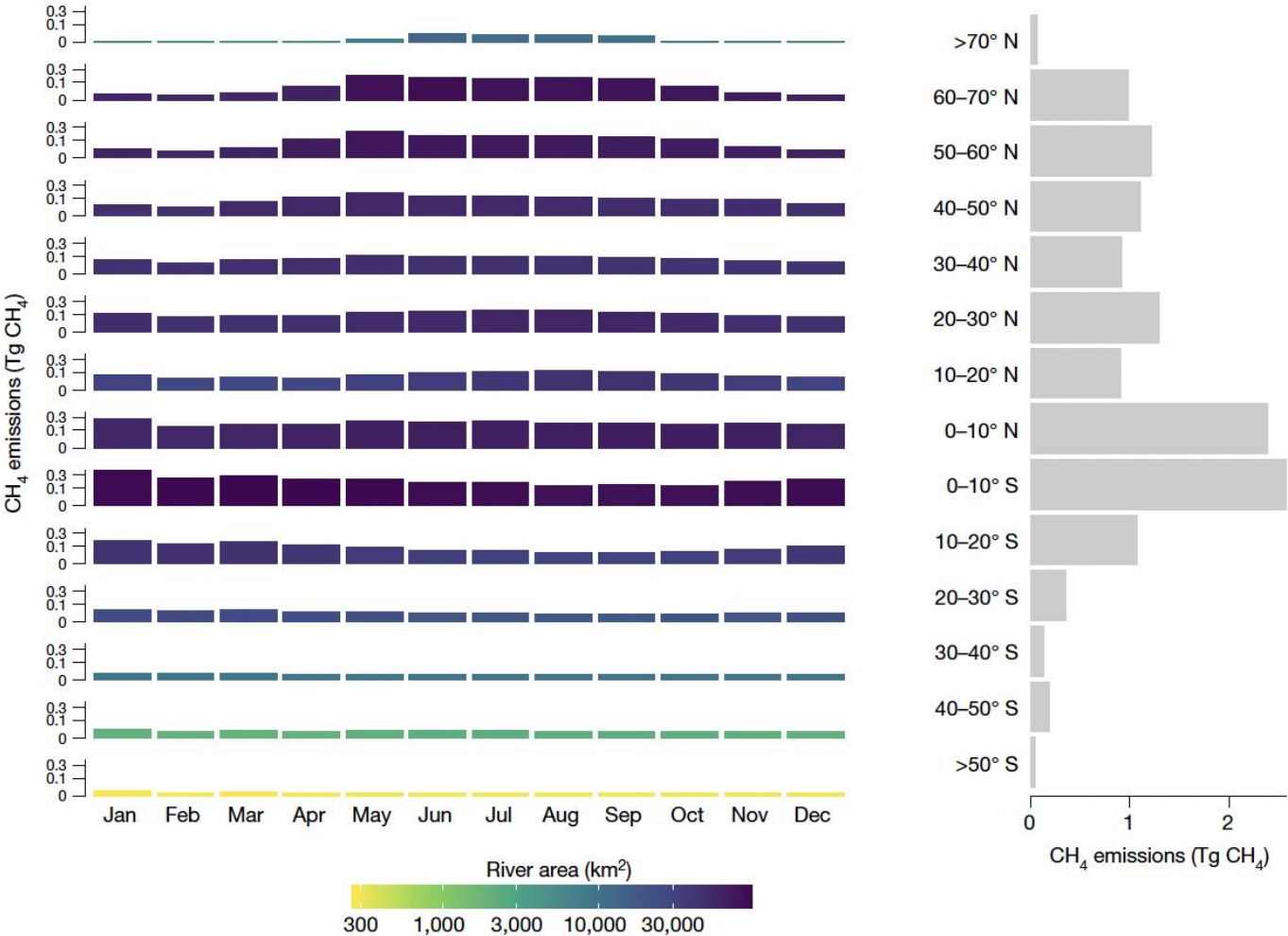
**Fig. 2 | Main drivers of CH<sub>4</sub> concentrations in streams. a,** The 20 most important variables in the random forest model. The x axis shows the median importance across all monthly models ( $n = 12$ ), with error lines representing standard deviation (s.d.); note the square-root transformation of the x axis. The line inside each bar is the partial dependence, which represents the marginal effect of a given feature (x axis) on predicted CH<sub>4</sub> concentrations (y axis). These lines are a simplification of a more detailed version (Supplementary Information). **b,** CH<sub>4</sub> concentrations of some site categories from GRiMeDB<sup>13</sup> were excluded from the model as they were not captured in the hydrological model or were targeted observations not representative of catchment properties (Methods). The underlying jittered points represent all other observations in GRiMeDB, with the dashed line representing the average. Each category is colour-coded, with the black dot and a line representing the mean  $\pm$  s.d.



Article

# Global methane emissions from rivers and streams

<https://doi.org/10.1038/s41586-023-06344-6> Gerard Rocher-Rog<sup>1,2,3,5</sup>, Emily H. Stanley<sup>4</sup>, Luke C. Loken<sup>5</sup>, Nora J. Casson<sup>6</sup>, Peter A. Raymond<sup>7</sup>, Shaoda Liu<sup>7</sup>, Giuseppe Amatulli<sup>7</sup> & Ryan A. Sponseller<sup>1</sup>  
Received: 25 October 2022

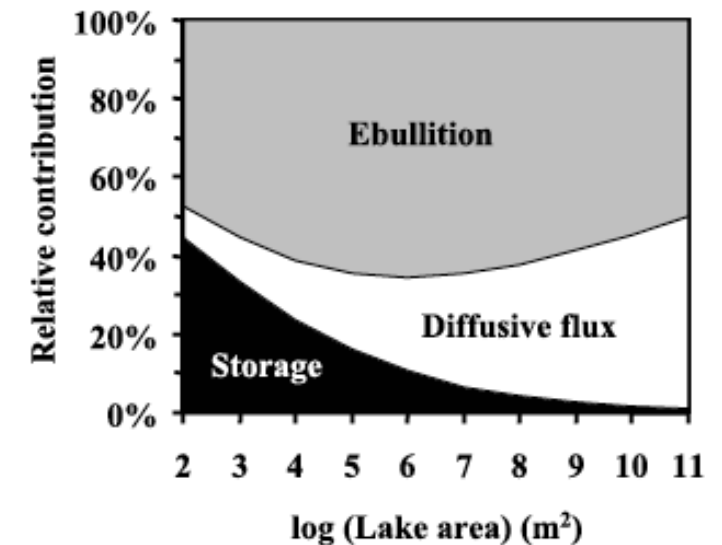
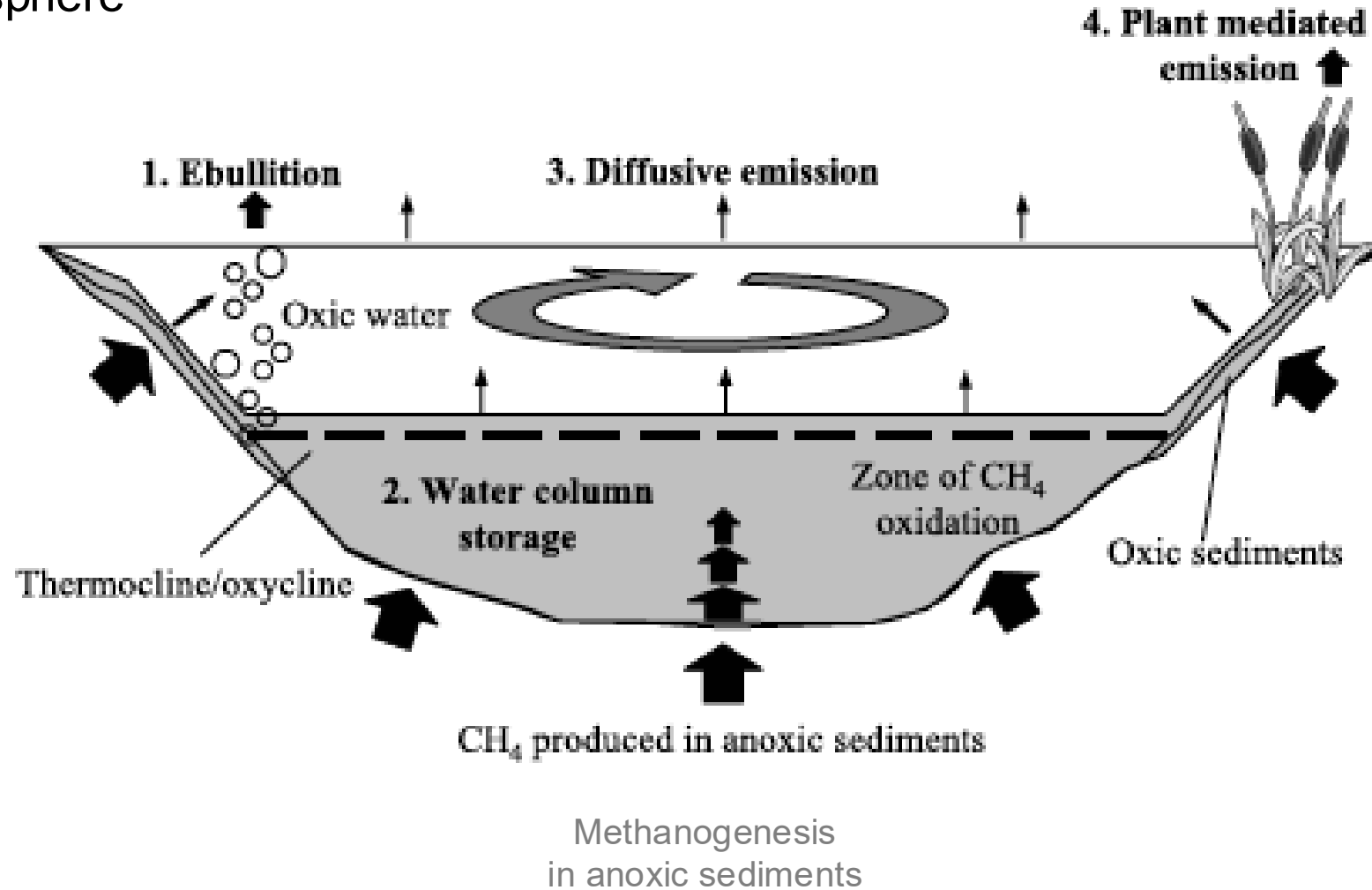


**Fig. 3 | Seasonal patterns of CH<sub>4</sub> emissions.** Left: total monthly CH<sub>4</sub> emissions for each latitudinal band (10° bins), with the colour representing total river area. Right: total yearly emissions for each latitudinal band. In the left panel, the y axis is square-root transformed, and the colour scale is log transformed.

- Clear seasonal and latitudinal patterns of CH<sub>4</sub> emissions
- Important to account for temporal and spatial variation of CH<sub>4</sub> emissions when scaling up to the global scale

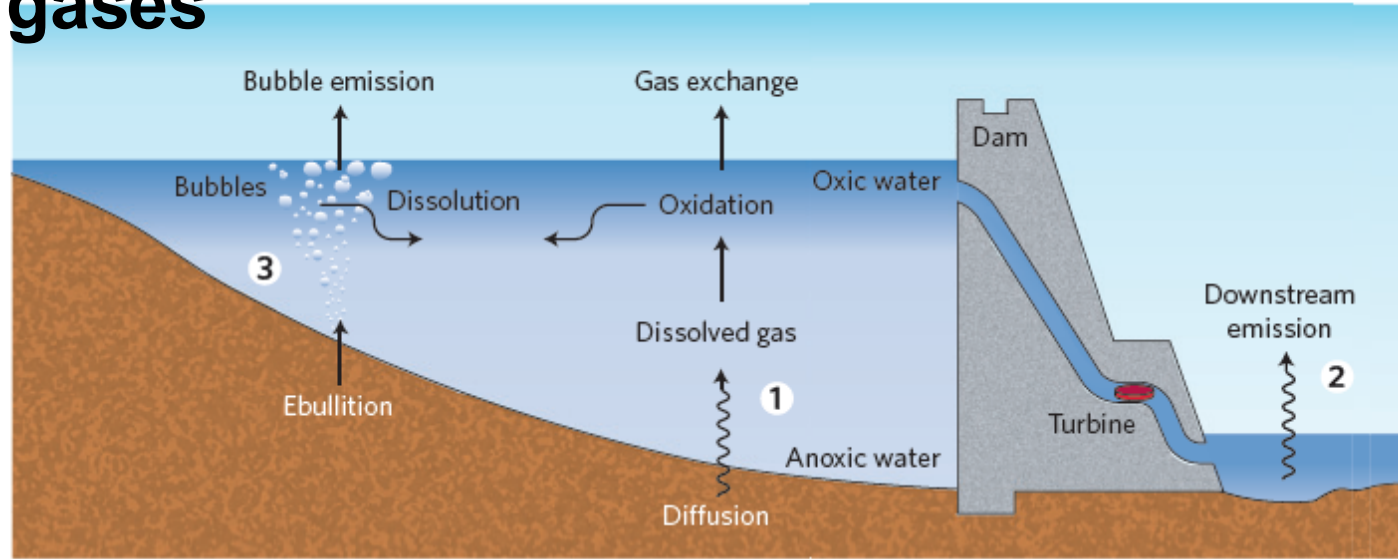
# Transportation of methane from aquatic ecosystems

Ebullition is the major transportation path from sediments to the water and atmosphere



Bastviken et al. 2004 GBC

# Hydropower reservoirs emit greenhouse gases



**Figure 1 |** Schematic methane emission pathways from a hydroelectric reservoir. In sediments with slower methane formation and at greater depths, dissolved methane diffuses upwards (1). The methane enters the atmosphere through gas exchange at the surface. Emissions may be reduced by microbial oxidation at the interface between the oxic and anoxic water layers. Second, downstream emissions after the water has passed the turbine (2) depend on the stratification of the reservoir and the vertical position of the main water intake. Finally, in sediments with high methane production rate, bubbles form when the methane solubility is exceeded (3). Some of this methane dissolves from the rising bubbles but a large fraction is rapidly emitted to the atmosphere. Barros and colleagues<sup>3</sup> estimate emissions of carbon dioxide and methane from 85 hydroelectric reservoirs worldwide, but because the third pathway is poorly constrained by measurements, uncertainties remain large.

# The contribution of dams to CH<sub>4</sub> production and outgassing

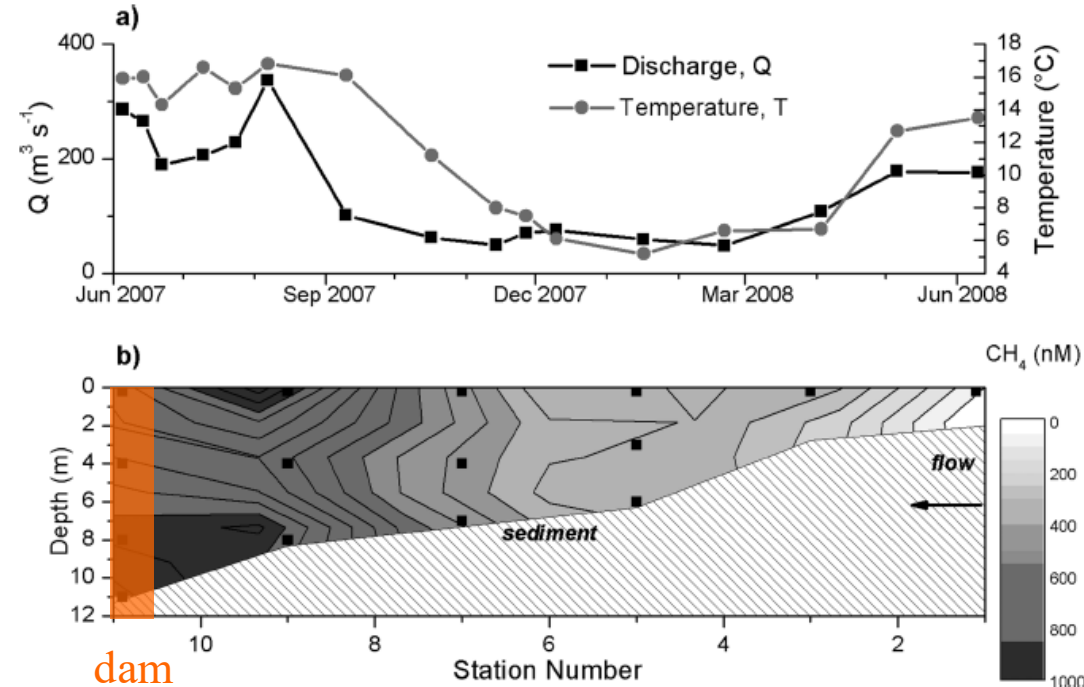


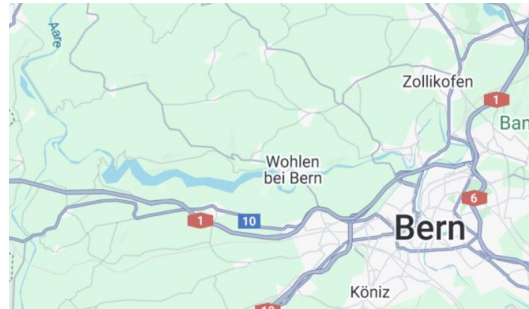
FIGURE 1. a) Temperature and discharge,  $Q$ , in Lake Wohlensee during samplings.  $Q$  ranges from  $\sim 400 \text{ m}^3 \text{s}^{-1}$  in summer (residence time,  $R_t$ ,  $\sim 1$  day) to  $50 \text{ m}^3 \text{s}^{-1}$  in winter ( $R_t \sim 7$  days). b) Contour plot of dissolved methane distribution in Lake Wohlensee on June 21, 2007. Black squares - actual samplings. Water flows in from right to the dam (left). Profiles were taken every kilometer at a vertical resolution of 3–4 m. The figure suggests that methane is vertically homogeneous, while concentrations can increase five times horizontally.

Del Sontro et al. 2010 ES&T

CH<sub>4</sub> can accumulate in reservoirs towards the dam



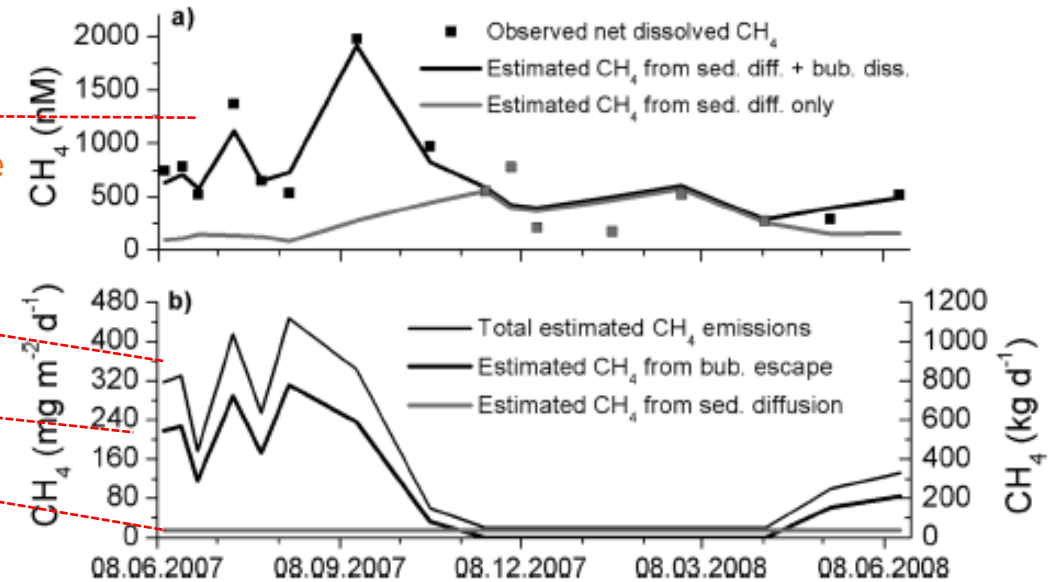
# The contribution of dams to CH<sub>4</sub> production and outgassing



Del Sontro et al. 2010 ES&T

Model including  
ebullition & temperature

total  
ebullition  
diffusion

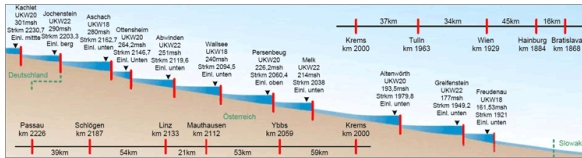


**FIGURE 3. Lake Wohlensee system analysis results. a)** Squares indicate measured dissolved methane in Lake Wohlensee (out-flow - inflow). Black squares indicate measurements at water temperatures,  $T > 10\text{ }^{\circ}\text{C}$  when ebullition is occurring, gray squares when  $T < 10\text{ }^{\circ}\text{C}$  (see Figure 2). Gray line - best fit of predicted concentration due to sediment diffusion (sed. diff.) only. Black line - model results for predicted CH<sub>4</sub> concentration using exponential fit for methane bubble dissolution (bub. diss.) as a function of temperature plus the constant sediment diffusion of  $15\text{ mg m}^{-2}\text{ day}^{-1}$  (Figure 2). **b)** CH<sub>4</sub> emission rates: Gray line - constant sediment diffusion input. Thick black line - predicted methane emission due to methane bubbles reaching the atmosphere (bub. escape). Thin black line - total predicted methane flux including dam discharge emissions.

Evasion flux:  $150\text{ mg CH}_4\text{ m}^{-2}\text{ d}^{-1}$

- Ebullition dominates the CH<sub>4</sub> evasion from Lake Wohlensee reservoir.
- Indicative of high CH<sub>4</sub> production rates in the accumulated sediments.





- Even low-land rivers with hydraulic constructions are sources of  $\text{CH}_4$
- Sediment accumulation and elevated  $\text{CH}_4$  production upriver from watergates for navigation

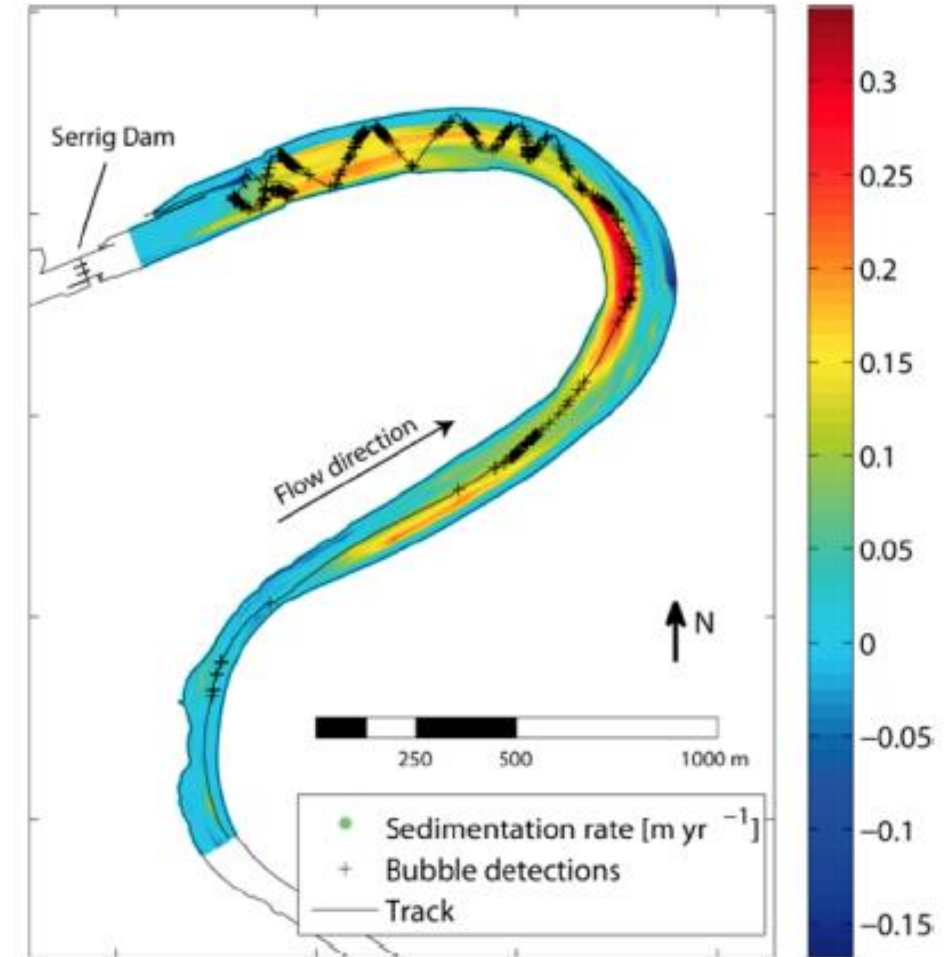


Figure 2. Mean sediment accumulation rate between 1993 and 2010 (color scaling) and bubble detections in the forebay of Serrig dam. Crosses mark bubble detection along the sampling transect, which is denoted by the black line.

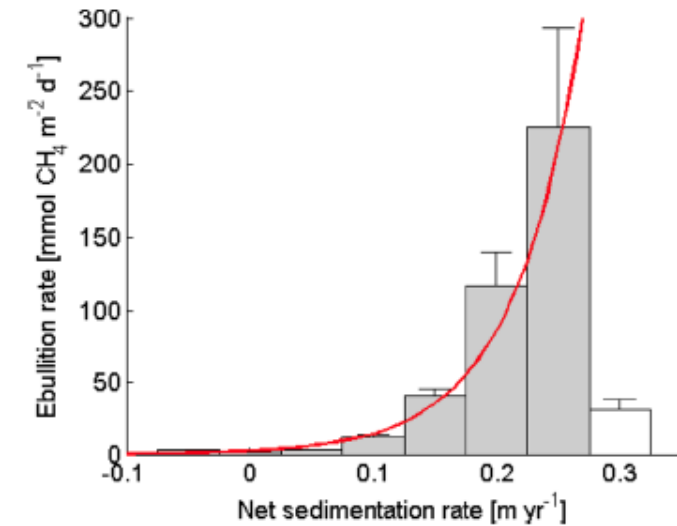
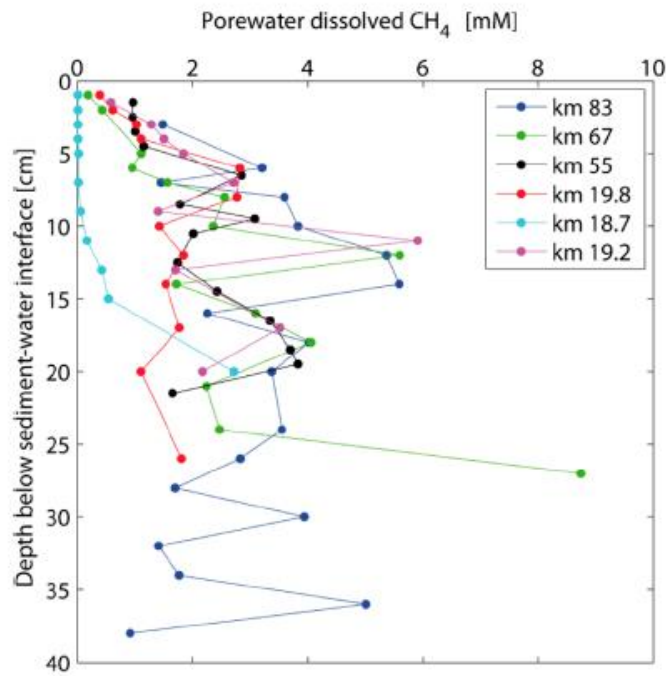


Figure 4. Relationship between sediment accumulation rate (1993–2010) and measured ebullition rates. The red line shows the exponential fit ( $R^2 = 0.91$ ;  $p < 0.001$ ;  $n = 7$ ). The white bar at 0.3 was excluded from the analysis due to its small sample size. Error bars denote the standard error of mean.

Maeck et al. 2013 *ES&T*

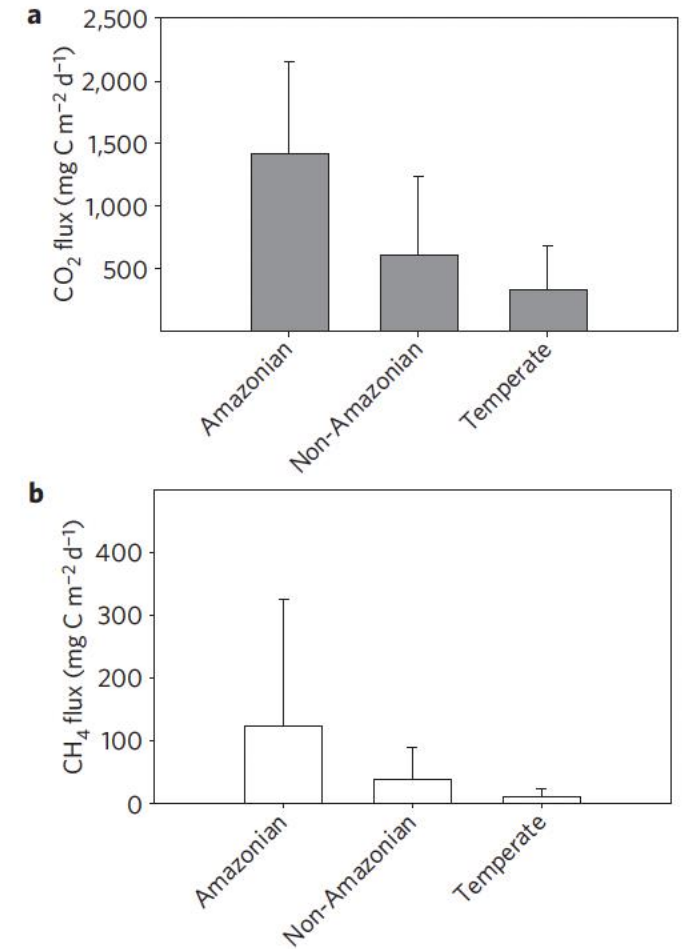
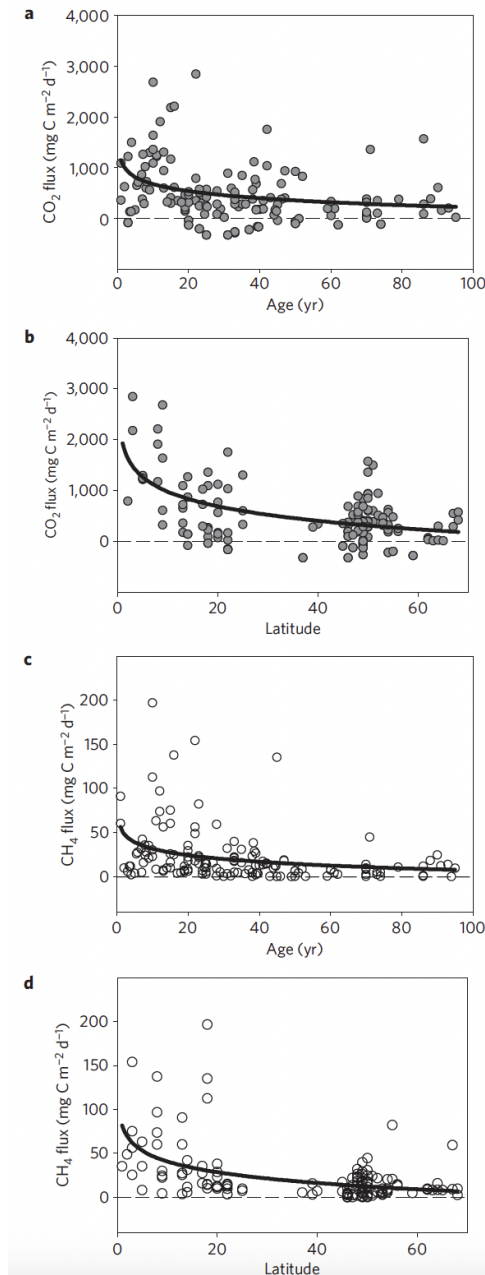
Reaches with elevated sediment accumulation are hotspots of  $\text{CH}_4$  production and ebullition in rivers

## Carbon emission from hydroelectric reservoirs linked to reservoir age and latitude

Nathan Barros<sup>1</sup>, Jonathan J. Cole<sup>2</sup>, Lars J. Tranvik<sup>3</sup>, Yves T. Prairie<sup>4</sup>, David Bastviken<sup>5</sup>, Vera L. M. Huszar<sup>6</sup>, Paul del Giorgio<sup>4</sup> and Fábio Roland<sup>1\*</sup>

- GHG emissions from reservoirs changing with latitude and age
- Highest in the Amazonas
- Topography, temperature and flooded biomass

Petit Saut Reservoir French Guyane



**Figure 2 | Fluxes of CO<sub>2</sub> and CH<sub>4</sub> in different zones.** Mean (bars) and standard deviation (lines) of the **a**, CO<sub>2</sub> and **b**, CH<sub>4</sub> fluxes in the 85 hydroelectric reservoirs worldwide distributed clustered by region. The tropical region was split into Amazonian and non-Amazonian regions.



# Greenhouse Gas Emissions from Reservoir Water Surfaces: A New Global Synthesis

BRIDGET R. DEEMER, JOHN A. HARRISON, SIYUE LI, JAKE J. BEAULIEU, TONYA DELSANTRO, NATHAN BARROS, JOSÉ F. BEZERRA-NETO, STEPHEN M. POWERS, MARCO A. DOS SANTOS, AND J. ARIE VONK

- Reservoirs are significant emitters of greenhouse gases ( $\text{CH}_4$ ,  $\text{CO}_2$  and  $\text{N}_2\text{O}$ ) to the atmosphere
  - Degradation of flooded biomass
  - Degradation of accumulating organic matter in the sediments
  - Increased temperature and primary production
- How carbon-free is hydropower?

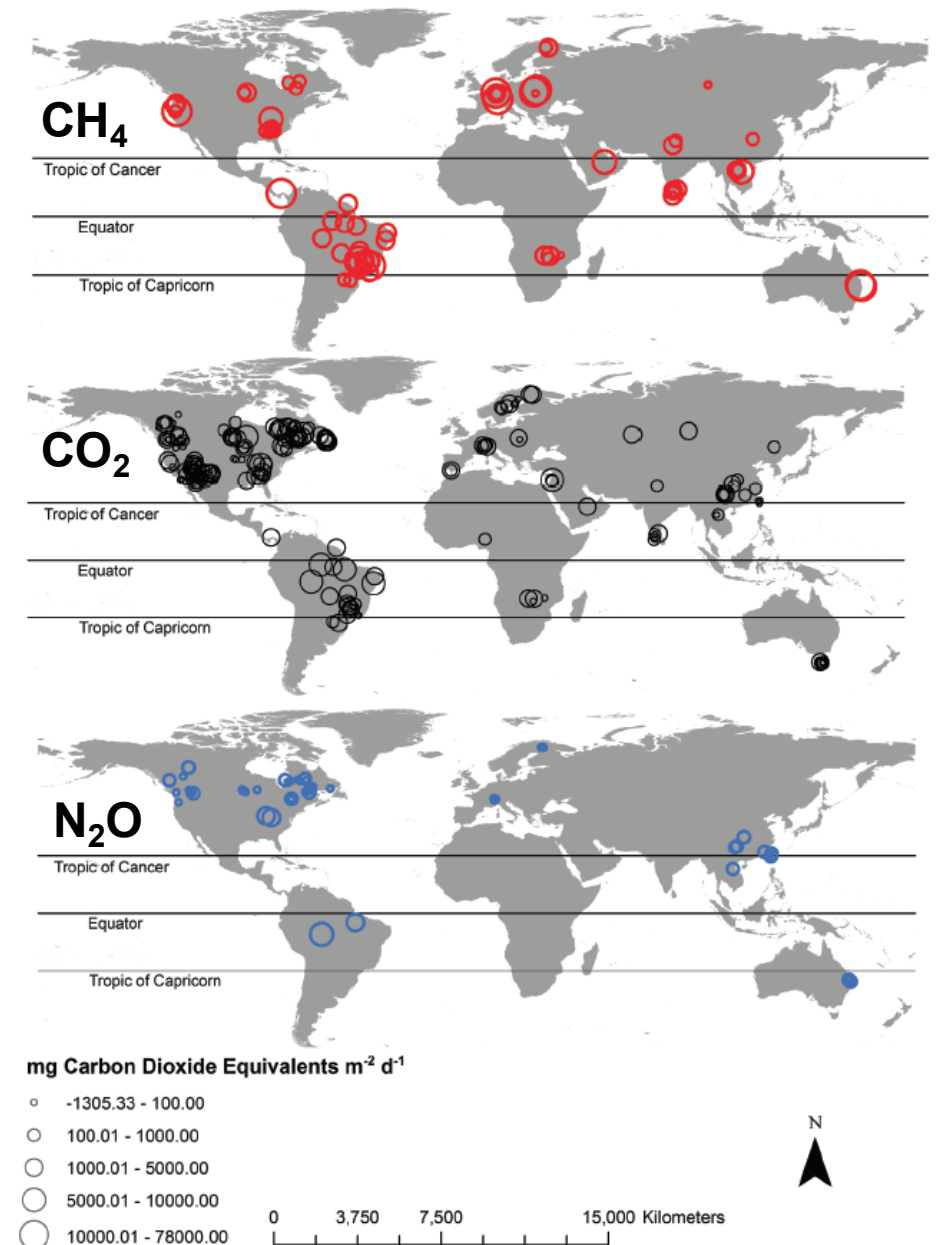


Figure 2. Diffusive + ebullitive methane (top), carbon dioxide (middle), and nitrous oxide (bottom) emissions from reservoirs on a  $\text{CO}_2$ -equivalent basis (100-year horizon). Few reservoirs had measurements for all three gases.

# Ecosystem size matters for GHG emissions

LETTERS

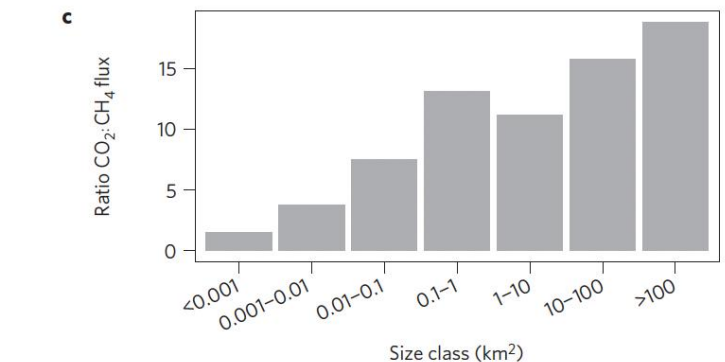
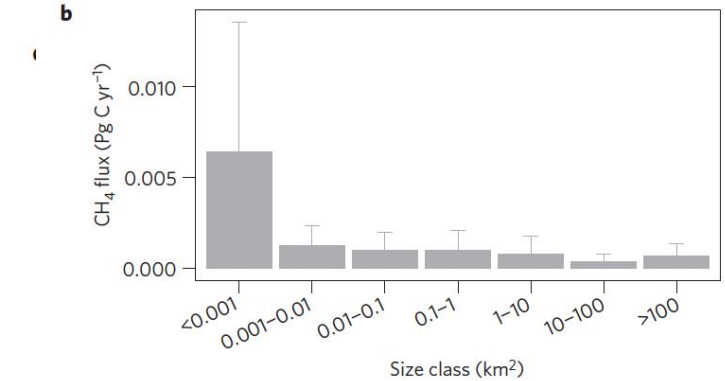
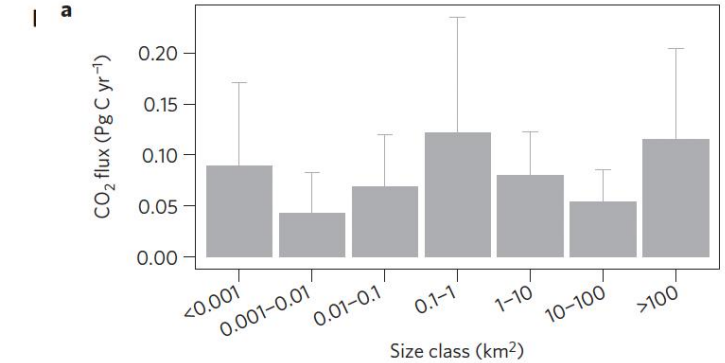
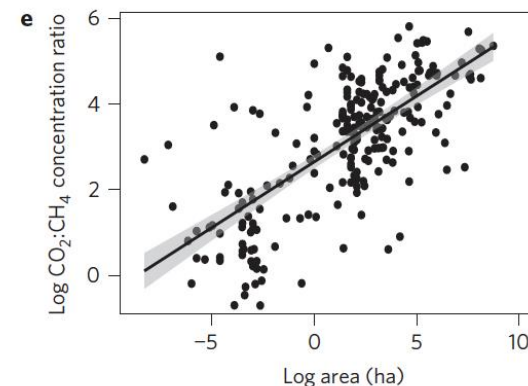
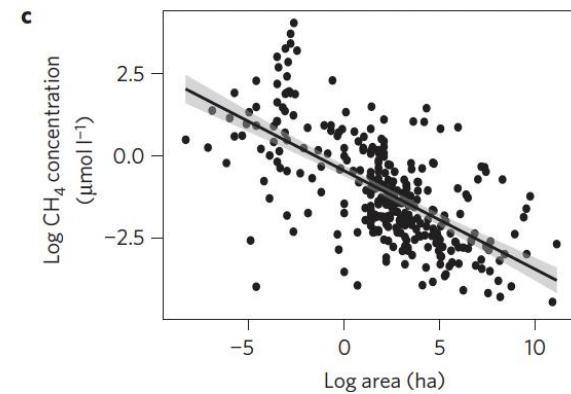
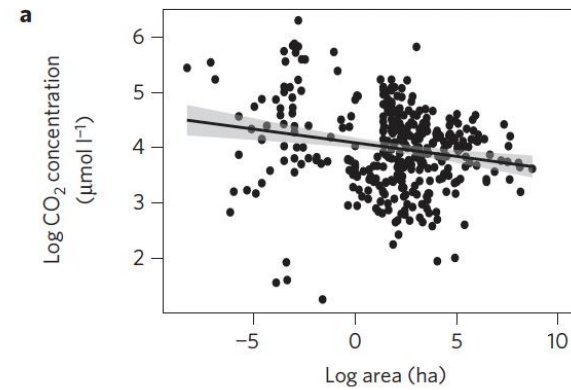
PUBLISHED ONLINE: 1 FEBRUARY 2016 | DOI: 10.1038/NGEO2654

nature  
geoscience

## Large contribution to inland water CO<sub>2</sub> and CH<sub>4</sub> emissions from very small ponds

Meredith A. Holgerson\* and Peter A. Raymond

- Small ponds emit more CH<sub>4</sub> than large ponds - also relative to CO<sub>2</sub> emissions
- Large receivers of organic matter and nutrients from the surrounding landscape (geometry effect)
- Organic matter burial and decomposition leading to hypoxic sedimentary habitats that promote methanogenesis



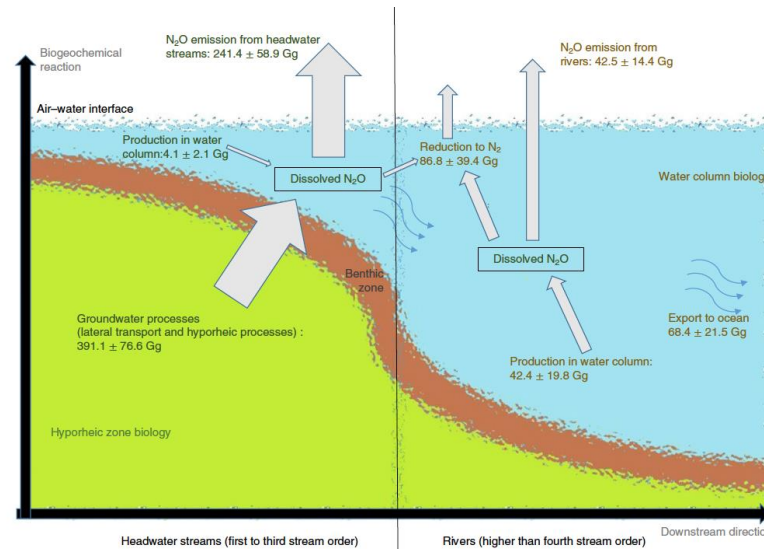


**Nitrous oxide ( $\text{N}_2\text{O}$ )**

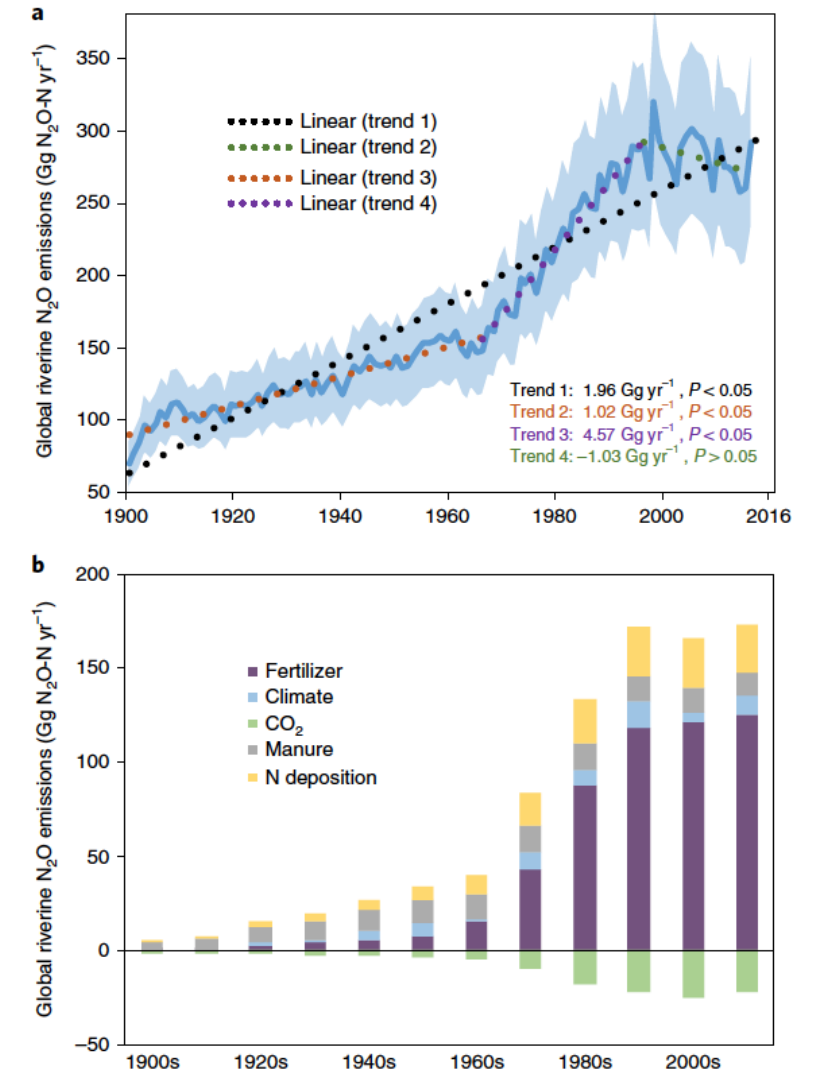
# Increased global nitrous oxide emissions from streams and rivers in the Anthropocene

Yuanzhi Yao<sup>1</sup>, Hanqin Tian<sup>1\*</sup>, Hao Shi<sup>1</sup>, Shufen Pan<sup>1</sup>, Rongting Xu<sup>1</sup>, Naiqing Pan<sup>1</sup> and Josep G. Canadell<sup>1,2</sup>

- $\text{N}_2\text{O}$  emissions from streams and rivers are increasing
- Major contributions from fertilizers (N rich), manure and atmospheric depositions
- Headwaters contribute most to  $\text{N}_2\text{O}$  emissions
- Tightly connected via shallow groundwater to land (with agricultural practices)



**Fig. 2 |** Global annual mean riverine  $\text{N}_2\text{O}$  fluxes during the 2000s estimated by DLEM. All the arrows denote  $\text{N}_2\text{O}$  fluxes. The left side of the figure depicts biogeochemical processes in the headwater zone simulated in subgrid routine processes at a resolution of  $0.5^\circ \times 0.5^\circ$ . The dissolved  $\text{N}_2\text{O}$  of headwater zone exports to downstream river channels (right side) were simulated through the DLEM cell-to-cell routine processes. The benthic zone indicates the sediment surface and its subsurface layers located at the lower end of the waterbodies.



**Fig. 1 |** Temporal pattern of global riverine  $\text{N}_2\text{O}$  emission and factorial analysis from 1900 to 2016. **a**, Global riverine  $\text{N}_2\text{O}$  emissions from 1900 to 2016 with uncertainty ranges shaded in blue ( $\pm 1$  s.d.). **b**, The factorial contributions to global riverine  $\text{N}_2\text{O}$  emissions from the 1900s to the period 2007–2016.

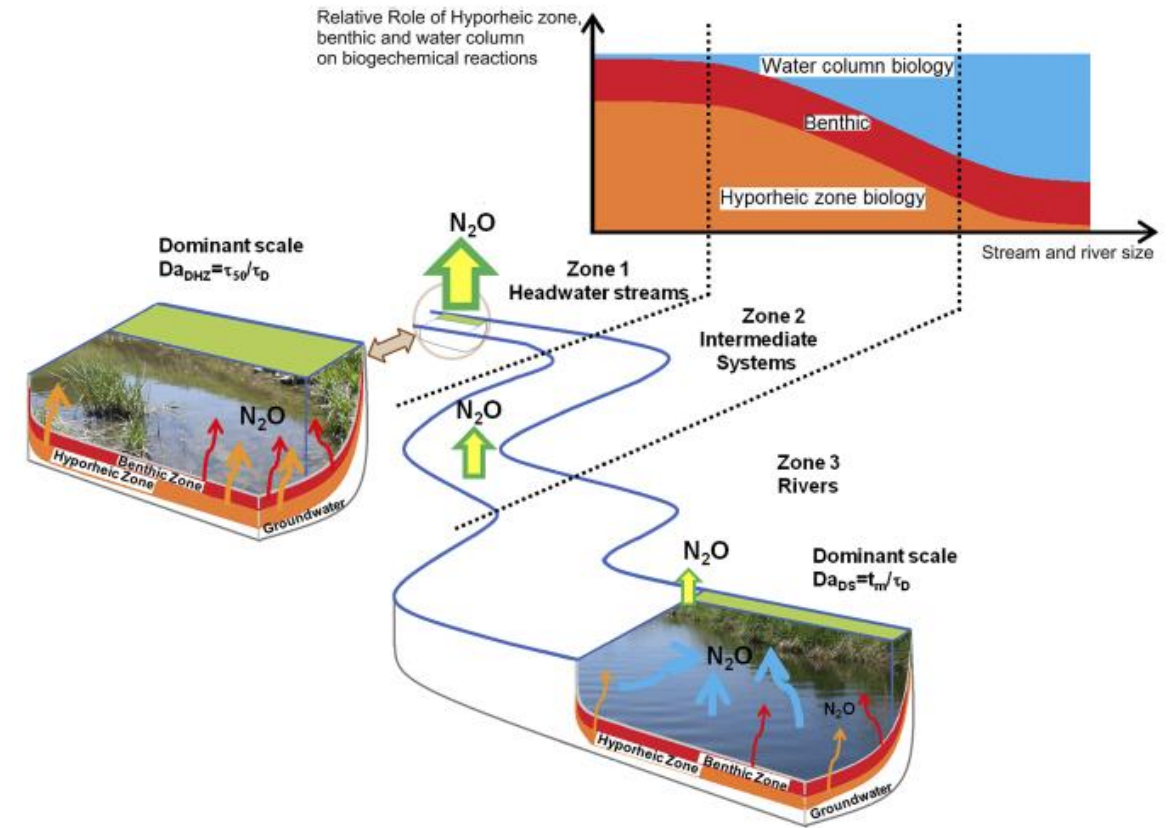
# Role of surface and subsurface processes in scaling $\text{N}_2\text{O}$ emissions along riverine networks

Alessandra Marzadri<sup>a,1,2</sup>, Martha M. Dee<sup>b,1</sup>, Daniele Tonina<sup>a,1</sup>, Alberto Bellin<sup>c,1</sup>, and Jennifer L. Tank<sup>b,1</sup>

Damköhler number for the benthic–hyporheic zone is defined as the ratio between the median hyporheic residence time ( $\tau_{50}$ ), which is an index of the time that streamwater spends within the hyporheic sediment, and the characteristic time of denitrification ( $\tau_D$ )

The dimension-less flux of  $\text{N}_2\text{O}$ ,  $F_{\text{N}_2\text{O}}$ , as the ratio between  $F_{\text{N}_2\text{O}}$  and the total flux per unit streambed area of dissolved inorganic nitrogen species [ $\text{NO}_3$  and  $\text{NH}_4$ ] in the stream ( $F_{\text{DINO}}$ ).

The time of turbulent vertical mixing,  $\tau_m$ , which is the average time for any neutrally buoyant particle to sweep through the water column because of turbulence. Damköhler number for rivers,  $\text{Da}_{\text{DS}} = \tau_m / \tau_D$ , with  $\tau_m$  replacing  $\tau_{50}$  and stating a shift from hyporheic to water column dominated  $\text{N}_2\text{O}$  production.



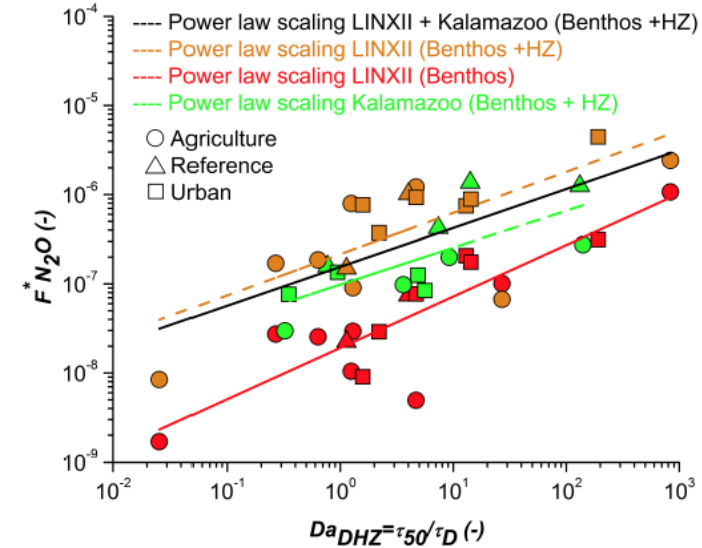
# Role of surface and subsurface processes in scaling $N_2O$ emissions along riverine networks

Alessandra Marzadri<sup>a,1,2</sup>, Martha M. Dee<sup>b,1</sup>, Daniele Tonina<sup>a,1</sup>, Alberto Bellin<sup>c,1</sup>, and Jennifer L. Tank<sup>b,1</sup>

Damköhler number for the benthic–hyporheic zone is defined as the ratio between the median hyporheic residence time ( $\tau_{50}$ ), which is an index of the time that streamwater spends within the hyporheic sediment, and the characteristic time of denitrification ( $\tau_D$ )

The dimension-less flux of  $N_2O$ ,  $F_{N_2O}$ , as the ratio between  $F_{N_2O}$  and the total flux per unit streambed area of dissolved inorganic nitrogen species [ $NO_3$  and  $NH_4$ ] in the stream ( $F_{DINO}$ ).

The time of turbulent vertical mixing,  $\tau_m$ , which is the average time for any neutrally buoyant particle to sweep through the water column because of turbulence. Damköhler number for rivers,  $Da_{DS} = \tau_m / \tau_D$ , with  $\tau_m$  replacing  $\tau_{50}$  and stating a shift from hyporheic to water column dominated  $N_2O$  production.



**Fig. 3.** Dimensionless flux of  $N_2O$  ( $F^*N_2O$ ) as a function of the denitrification Damköhler number ( $Da_{DHZ}$ ) in the LINXII Study ( $n$  = number of streams,  $n = 16$ ) and the Kalamazoo River (Michigan;  $n = 12$ ) streams.  $F^*N_2O$  resulting from the production of  $N_2O$  within only the benthic zone of the LINXII Study streams is shown with red symbols; the power law regression of these data is shown with the red solid line [ $F^*N_2O = 1.91 \times 10^{-8}(Da_{DHZ})^{0.57}$ ,  $r^2 = 0.75$ ]. Emissions from the benthic–hyporheic zone (combined contribution of both zones, Benthos + HZ) are in orange symbols, and their power regression is shown as the orange dashed line [ $F^*N_2O = 2.15 \times 10^{-7}(Da_{DHZ})^{0.46}$ ,  $r^2 = 0.54$ ]. Emissions from the benthic–hyporheic zone of the Kalamazoo streams scale with  $Da_{DHZ}$  [ $F^*N_2O = 9.83 \times 10^{-8}(Da_{DHZ})^{0.41}$ ,  $r^2 = 0.54$ ] as shown by the green line. Because these two relationships (dashed orange and green lines) are not significantly different, we fitted both datasets with a power law [ $F^*N_2O = 1.55 \times 10^{-7}(Da_{DHZ})^{0.43}$ ,  $r^2 = 0.48$ ; black line], which quantifies  $N_2O$  emissions from headwaters.





## Global riverine nitrous oxide emissions: The role of small streams and large rivers



Alessandra Marzadri <sup>a,\*</sup>, Giuseppe Amatulli <sup>b,c,\*</sup>, Daniele Tonina <sup>d</sup>, Alberto Bellin <sup>a</sup>, Longzhu Q. Shen <sup>e</sup>, George H. Allen <sup>f</sup>, Peter A. Raymond <sup>b</sup>

The application of this hybrid modelling approach reveals that small streams (width < 10 m) are the primary sources of riverine N<sub>2</sub>O emissions to the atmosphere.

They contribute nearly 36 Gg N<sub>2</sub>O–N/yr; almost 50% of the entire N<sub>2</sub>O emissions from riverine systems (72.8 Gg N<sub>2</sub>O–N/yr), although they account for only 13% of the total riverine surface area worldwide.

Large rivers (widths >175 m), such as the main stems of the Amazon River (~ 6 Gg N<sub>2</sub>O–N/yr), the Mississippi River (~ 2 Gg N<sub>2</sub>O–N/yr), the Congo River (~ 1 Gg N<sub>2</sub>O–N/yr) and the Yang Tze River (~ 0.7 Gg N<sub>2</sub>O–N/yr), only contribute 26% of global N<sub>2</sub>O emissions, which primarily originate from their water column.

Underscores the role of hyporheic processes in small streams for N<sub>2</sub>O production and emissions

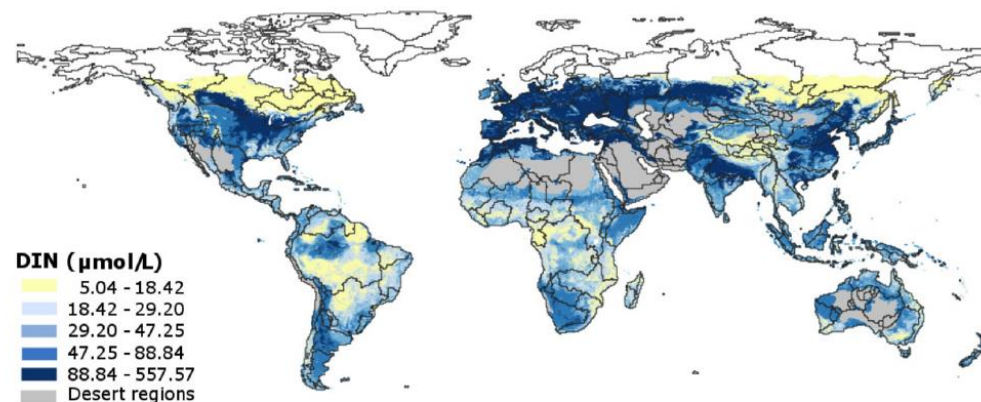


Fig. 1. Map of average annual Dissolved Inorganic Nitrogen (DIN) concentration distribution obtained by the data-driven (Random Forest) model (Shen et al., 2020). DIN map obtained as the combination of the in-stream load of ammonium and nitrate  $DIN = [NH_4^+] + [NO_3^-]$ , ( $\mu\text{mol/L}$ ).

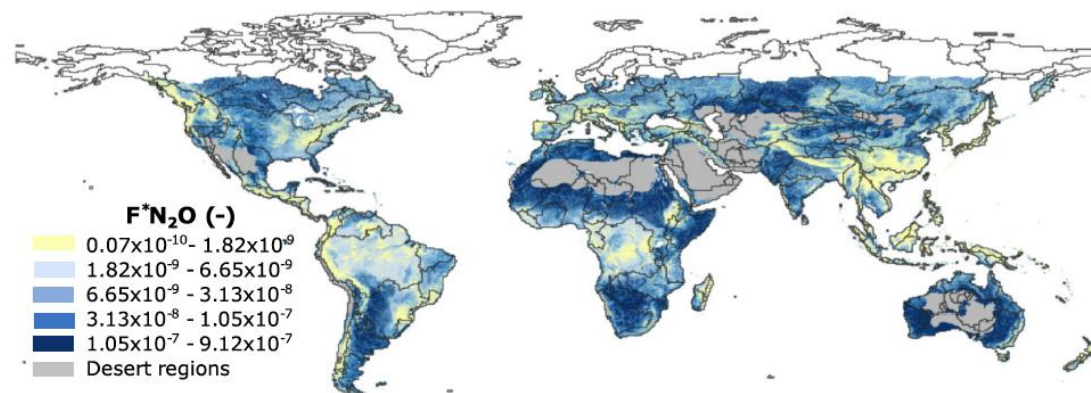


Fig. 4. Map of dimensionless N<sub>2</sub>O flux ( $F^*N_2O$ ) along the world river network analyzed. Gray areas represent desert regions not accounted in the calculation.



# Stream and river networks

## The multiple dimensions

### Vertical

- Connected to the atmosphere through the turbulent surface
- Connected to the groundwater

### Lateral

- Connected to groundwater, riparian zone and corridor

### Longitudinal

- Ample opportunities for downstream processing (see RCC)

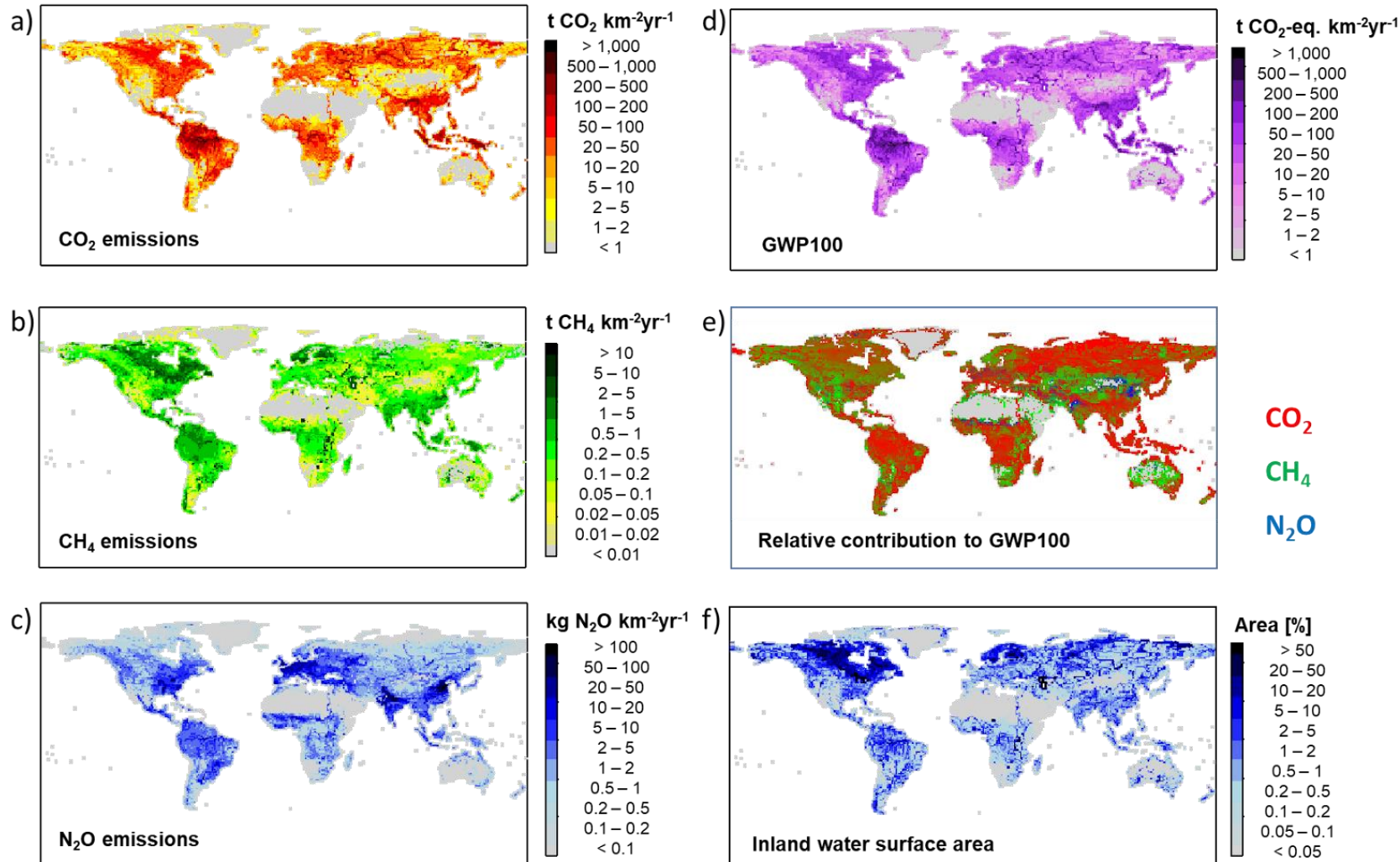
### Network

- Small streams are most abundant and tightly connected to the terrestrial environment

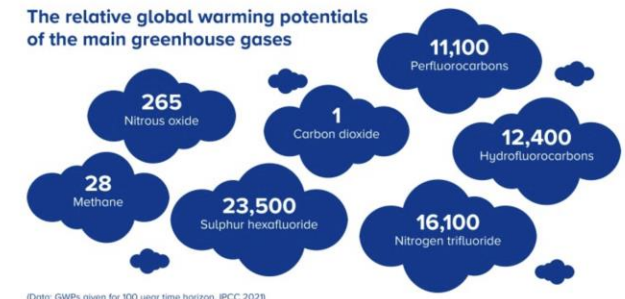
**Makes streams and rivers so important for carbon fluxes despite their minor contribution by areal extent**

# GHG emissions from inland waters

Global Warming Potential at 100 year horizon  
(GWP100)



- GWP100 of inland water GHG emissions amounts to ~7.6 (4.7-13.0) Pg CO<sub>2</sub>-eq yr<sup>-1</sup>.
- Roughly three quarters are contributed by net emissions of CO<sub>2</sub>, the remainder mainly CH<sub>4</sub>, while contributions of N<sub>2</sub>O emissions are nearly negligible.
- Overall streams and rivers emit ca. 80% of inland water GHG



(Data: GWPs given for 100 year time horizon, IPCC 2021)

Different GHGs have different global warming potentials

# Inland waters and the carbon cycle



Streams, rivers, lakes and ponds are critical components of the global carbon cycle

Tight connection with the terrestrial environment; they receive large terrestrial deliveries of organic matter and  $\text{CO}_2$  (from weathering and soil respiration)

Important sources of  $\text{CO}_2$ ,  $\text{CH}_4$  and  $\text{N}_2\text{O}$  to the atmosphere (uncertainties are large)

Inland waters are biogeochemical connectors between terrestrial ecosystems, atmosphere and the ocean