

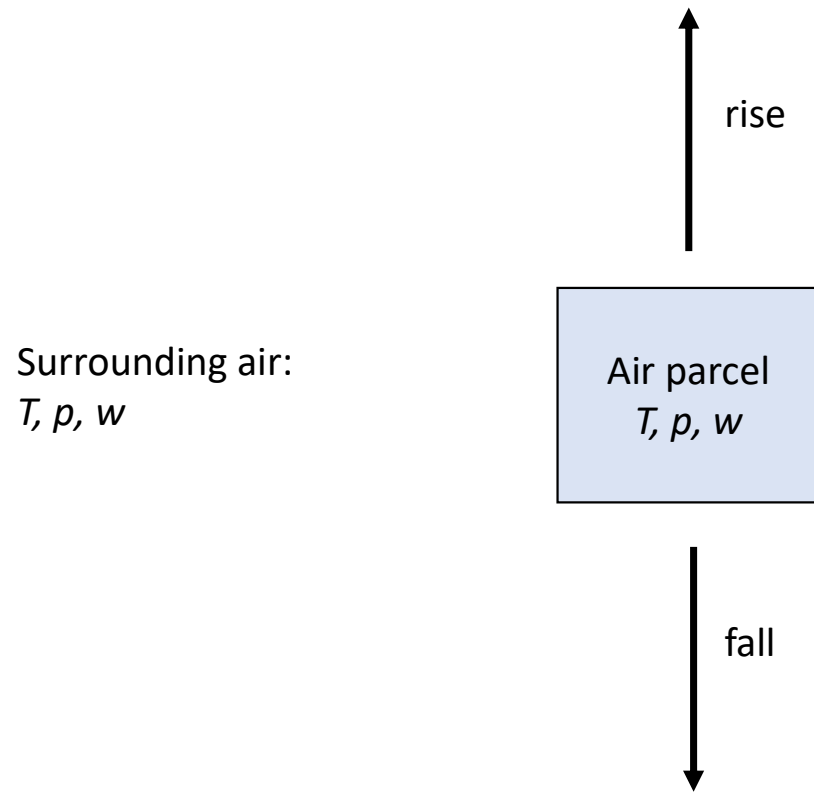
Physics and Chemistry of the Atmosphere

Assignment 1 - Introduction

Information

- Assignment can be found on Moodle
- Groups of 2 people maximum
- Submit solutions on Moodle
 - PDFs (handwritten solutions must be scanned)
 - Python file
- Provide details of your calculations for all answers
- Grading: $(\text{total points} / \text{total possible points}) \times 5 + 1$
- **Due date: Friday, 07.03.2025 23:59**

Recap: concept of an air parcel



Simplifying assumptions:

- No heat (i.e. energy) exchange with the environment \rightarrow *adiabatic*
- Surrounding air is still (i.e. fixed temperature gradient)

Diabatic vs. Adiabatic Processes

Diabatic process: Direct heat exchange with the environment.

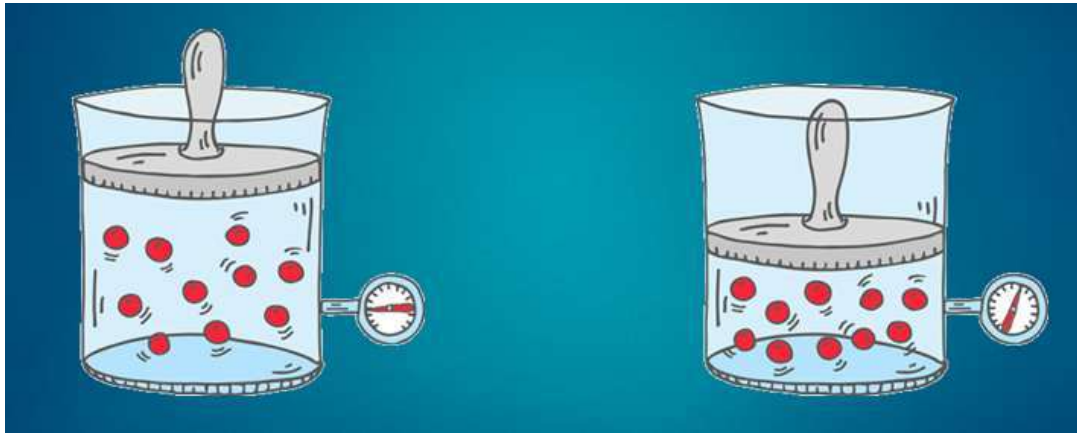
Example: heat exchange between air and surface



Diabatic vs. Adiabatic Processes

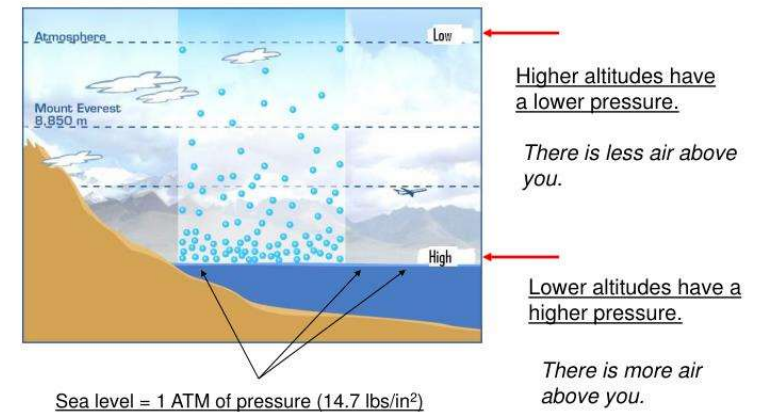
Adiabatic process: No heat exchange with the environment.

- *Heating due to compression*
- *Cooling due to expansion*



<https://www.air-compressor-guide.com/articles/water-in-compressed-air-calculations>

Atmospheric Pressure



<https://www.inspiritvr.com/air-pressure-at-mount-everest-study-guide/>

Recap: thermodynamics

1st law of thermodynamics:

$$dq = c_p dT - \frac{1}{\rho} dp;$$

$[q] = \text{J kg}^{-1}$
 $[T] = \text{K}$
 $c_p = 1004 \text{ J kg}^{-1} \text{ K}^{-1}$

heat exchange → change in internal energy → work

Adiabatic: no heat exchange

$$0 = c_p dT - \frac{1}{\rho} dp;$$

Change in temperature goes along with change in pressure

Recap: potential temperature

Potential temperature:

Temperature an air parcel would have when adiabatically brought to sea level (1000 hPa).

$$\theta = T \left(\frac{p_0}{p} \right)^{R/c_p}$$

$$R \approx 287 \text{ J kg}^{-1} \text{ K}^{-1}$$

$$c_p \approx 1000 \text{ J kg}^{-1} \text{ K}^{-1}$$

-> can also be read from skewT-logP diagram!

Recap: moisture in the atmosphere

Mixing ratio:

$$w = \frac{m_v}{m_d}$$

m_v : mass of water vapor in the air [kg] or [g]

m_d : mass of dry air [kg]

Saturation:

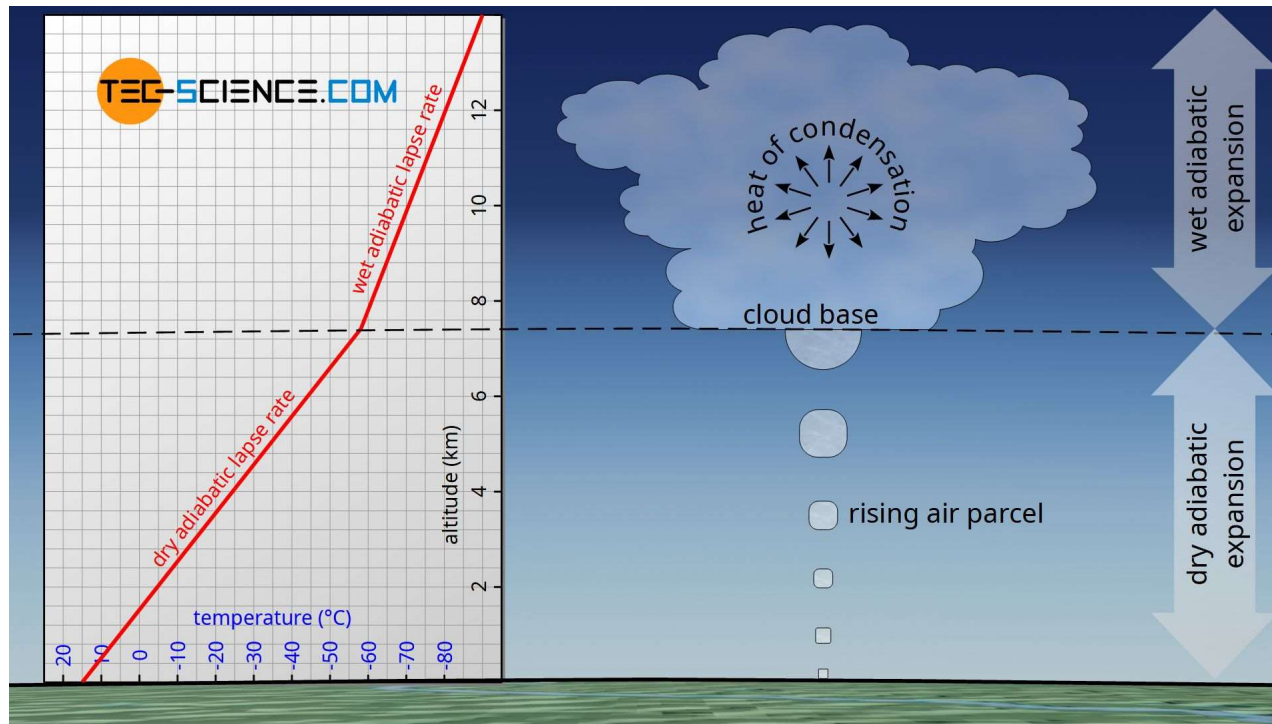
The maximum amount of water vapor an air parcel can contain at a specific temperature and pressure.

Relative humidity:

$$RH = \frac{w}{w_s} * 100\%$$

w_s : saturated mixing ratio [kg/kg] or [g/kg]

Adiabatic lapse rates



<https://www.tec-science.com/mechanics/gases-and-liquids/barometric-formula-for-an-adiabatic-atmosphere/>

Lifting Condensation Level (LCL)

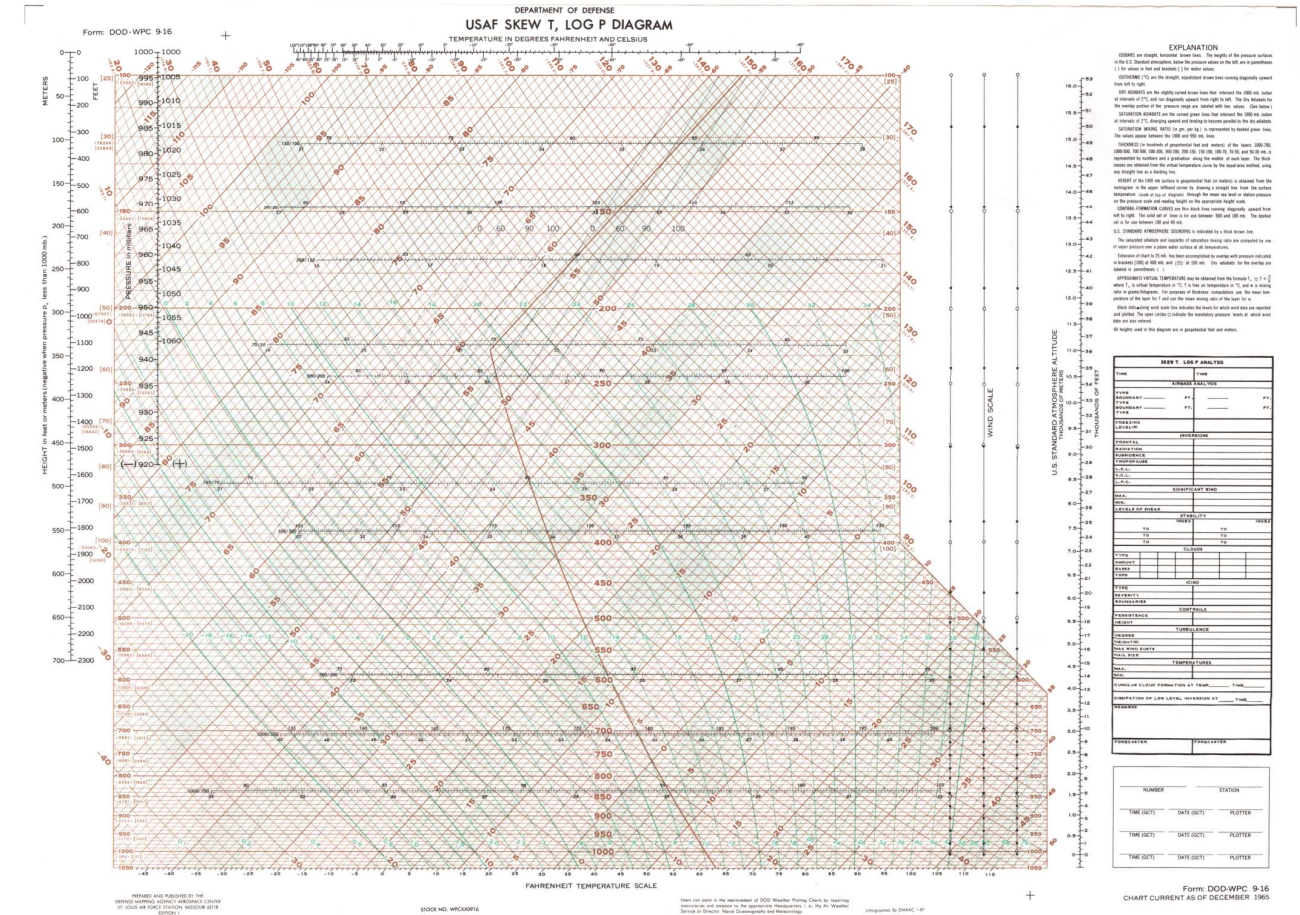
- Air is saturated.
- Further lifting of air parcel results in condensation (i.e. cloud forming).
- Condensation releases energy, that reduces the temperature gradient.

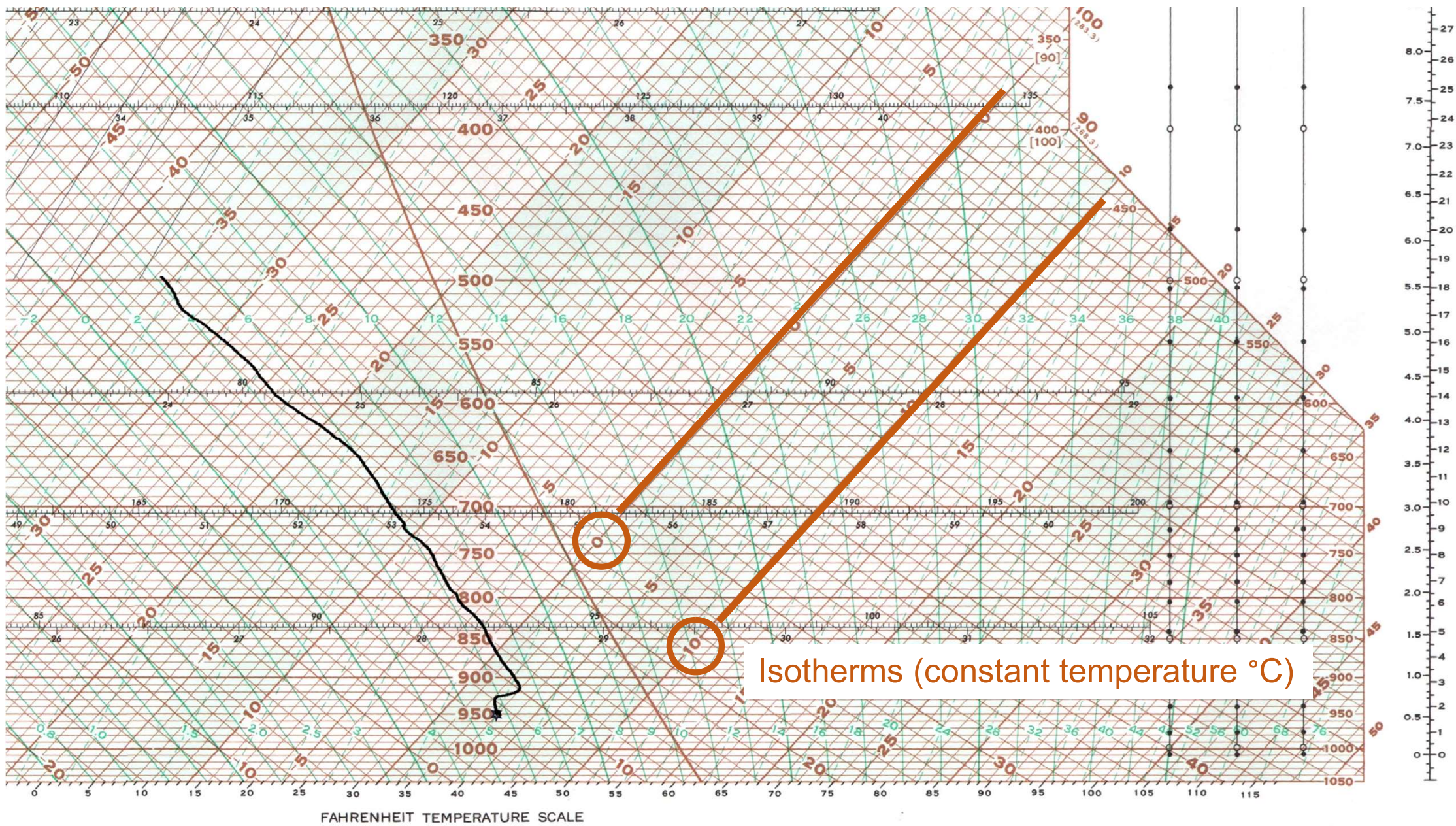
Wet/dry adiabatic lapse rate:

Temperature change of a lifted air parcel [°C/m]

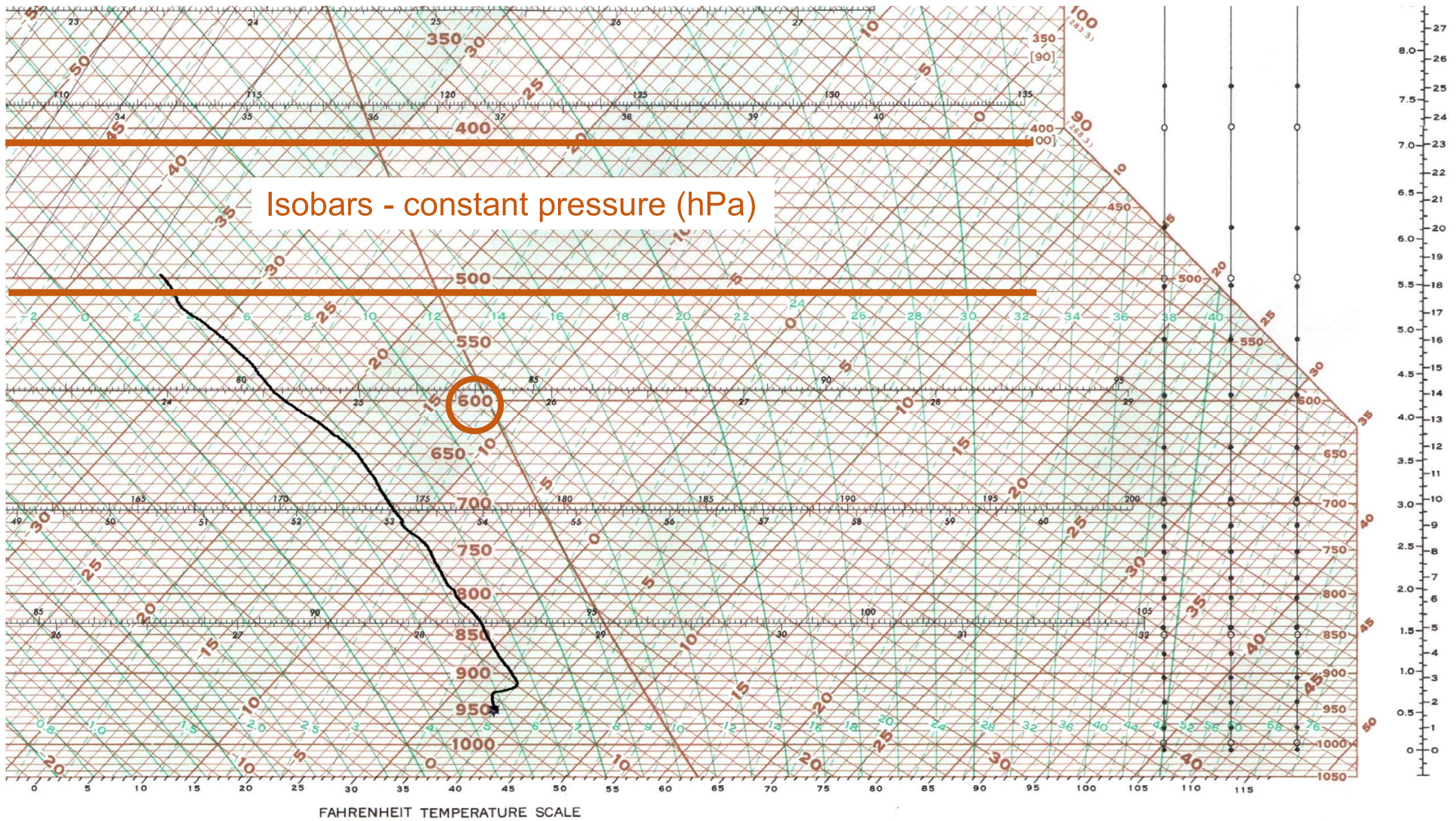
Environmental lapse rate:

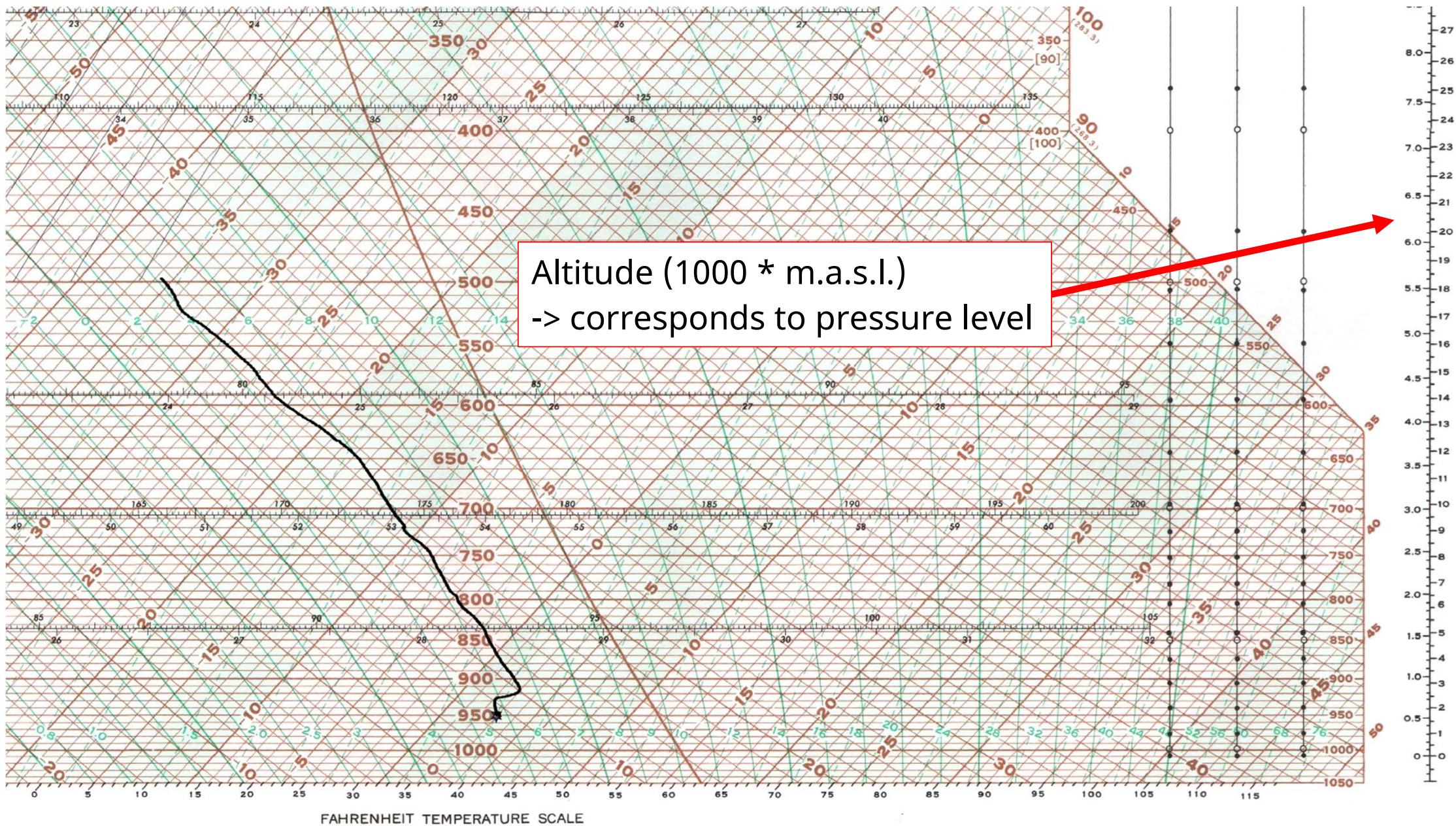
Temperature change of environment with altitude [°C/m]

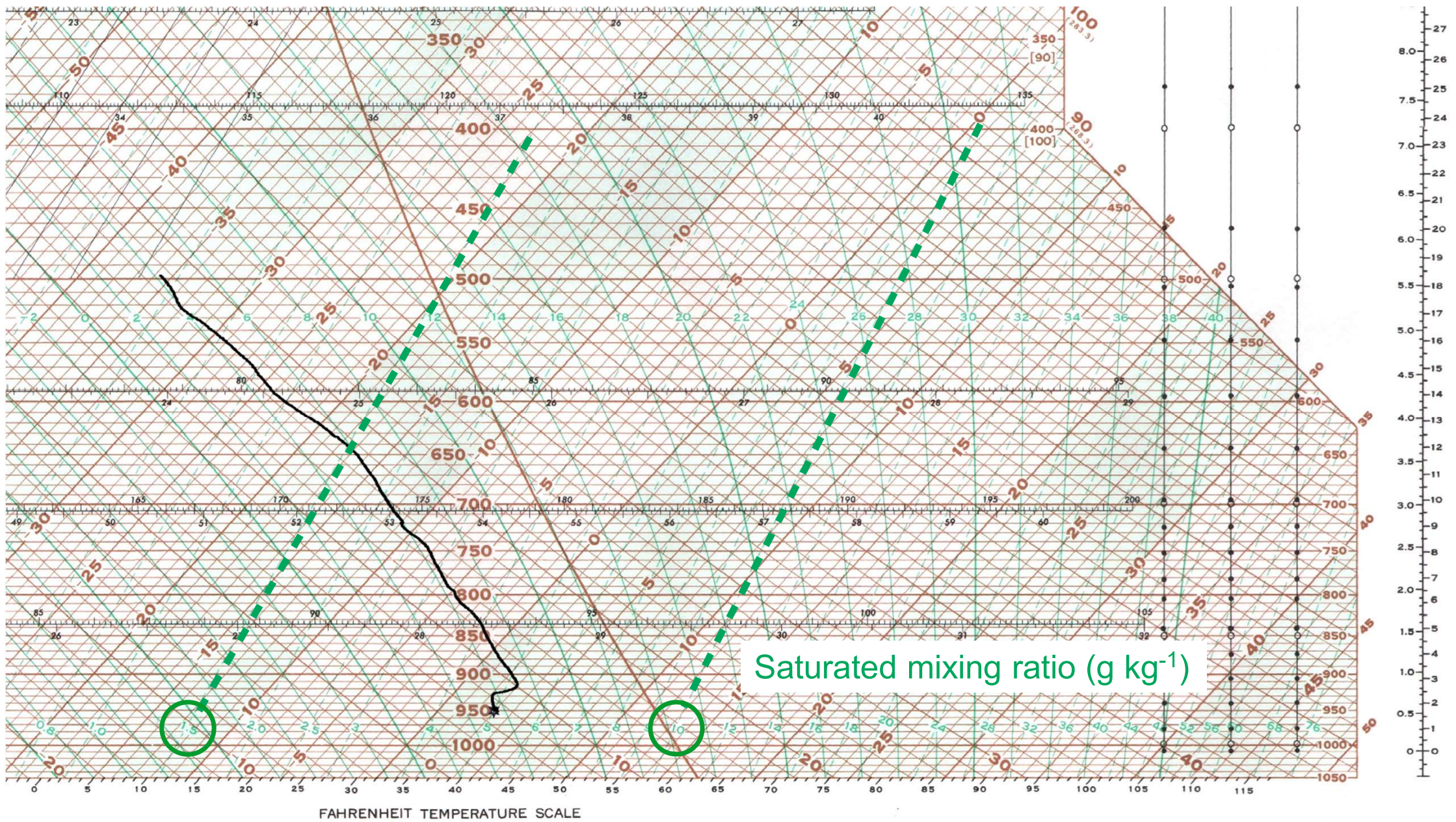


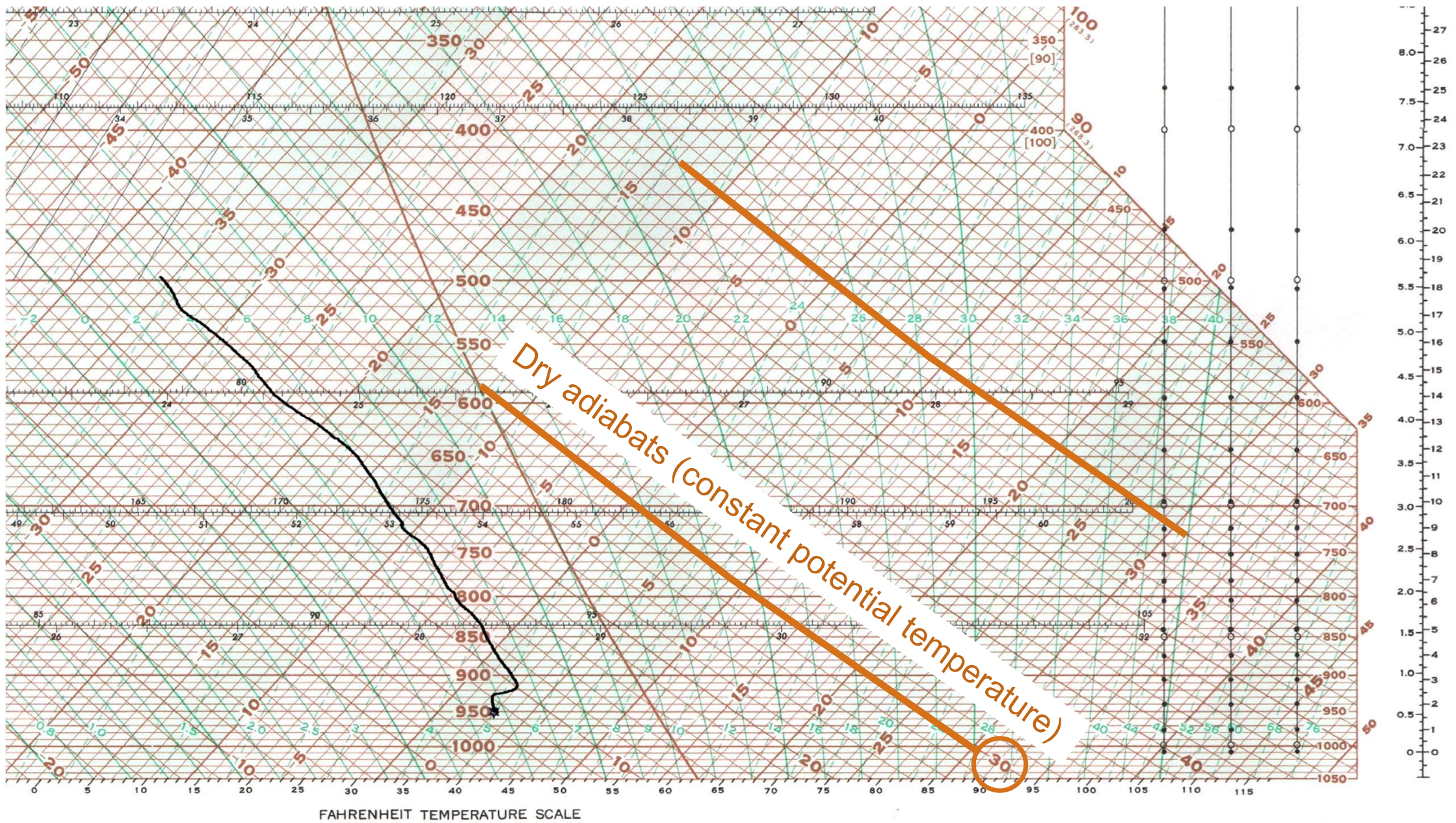


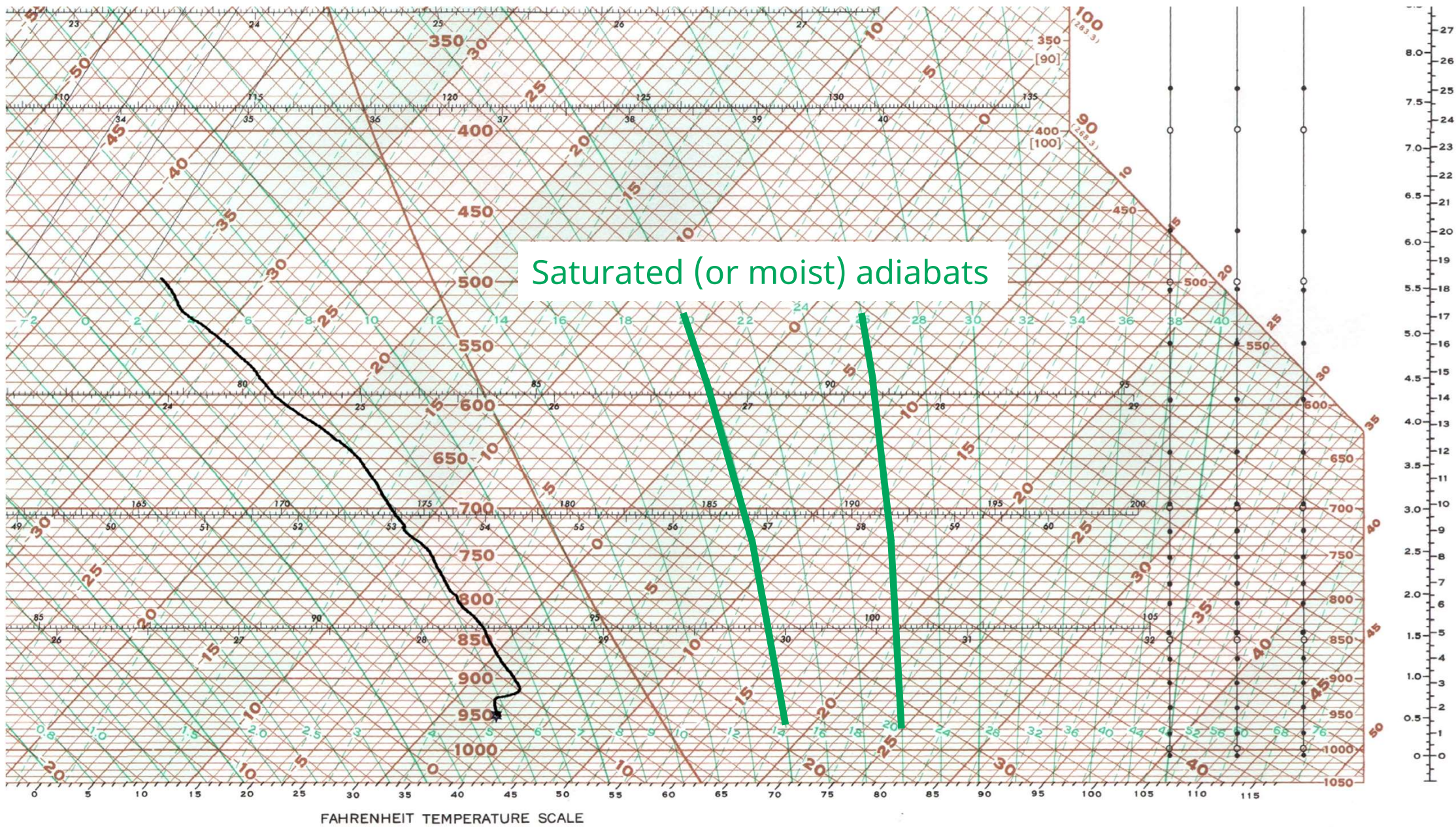
Isobars - constant pressure (hPa)

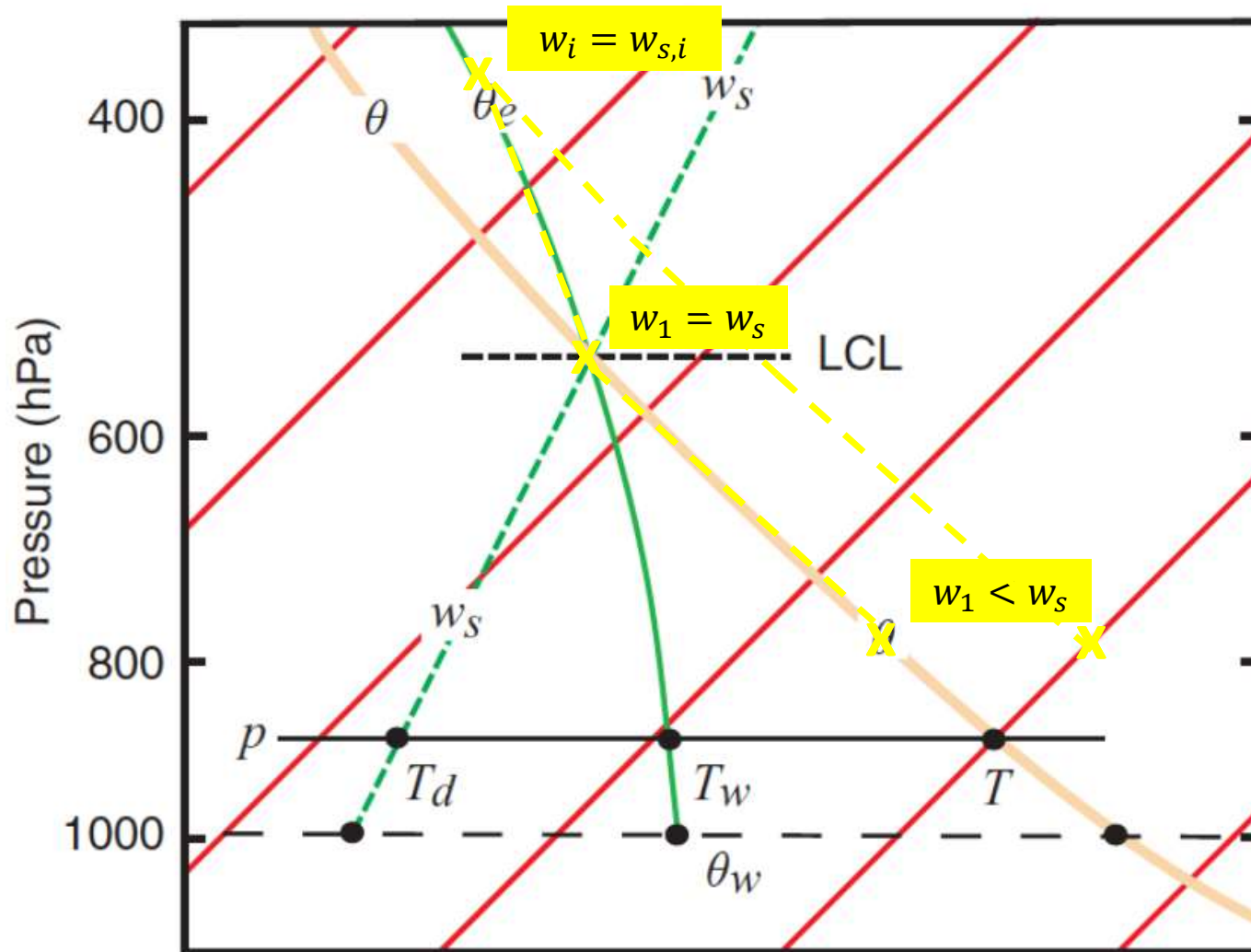












θ : potential temperature
 T : temperature
 T_w : wet bulb temperature
 T_d : dew point temperature
 w_s : saturated mixing ratio
 LCL: lifting condensation level

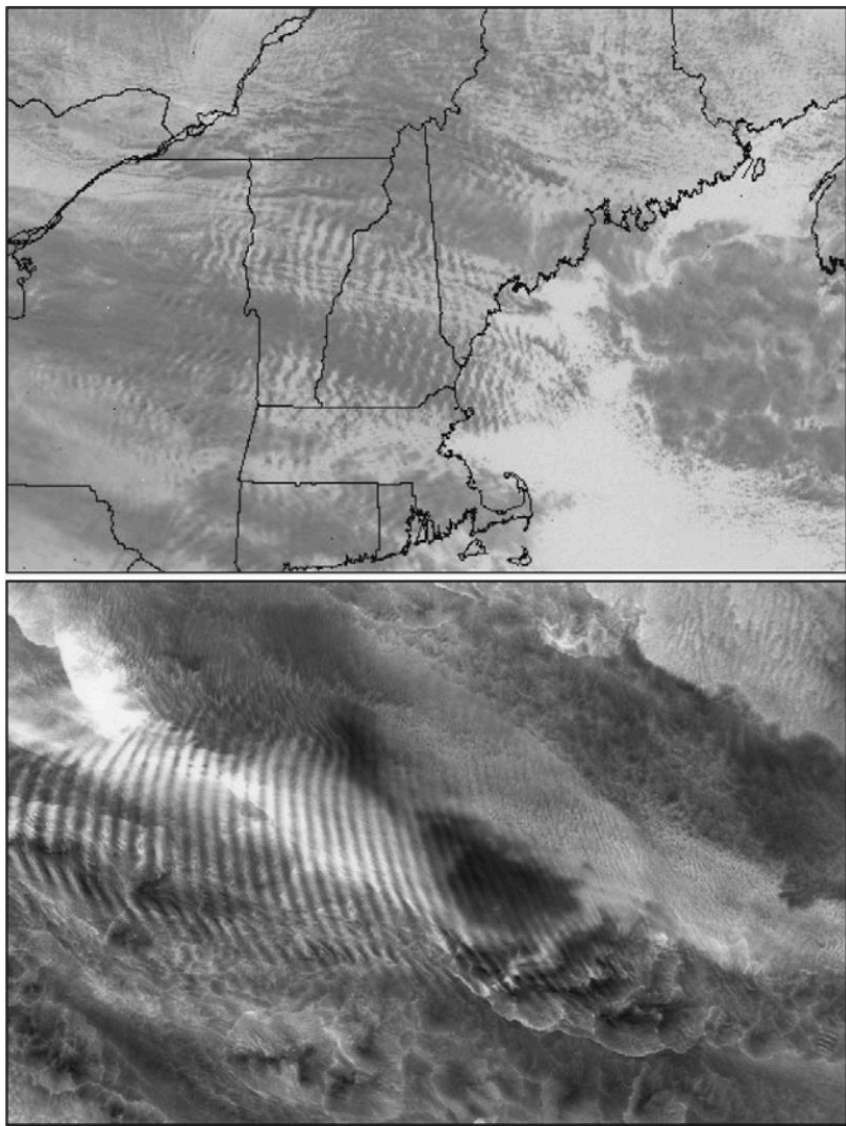


Fig. 3.14 Gravity waves, as revealed by cloud patterns. The upper photograph, based on NOAA GOES 8 visible satellite imagery, shows a wave pattern in west to east (right to left) airflow over the north-south-oriented mountain ranges of the Appalachians in the northeastern United States. The waves are transverse to the flow and their hori-

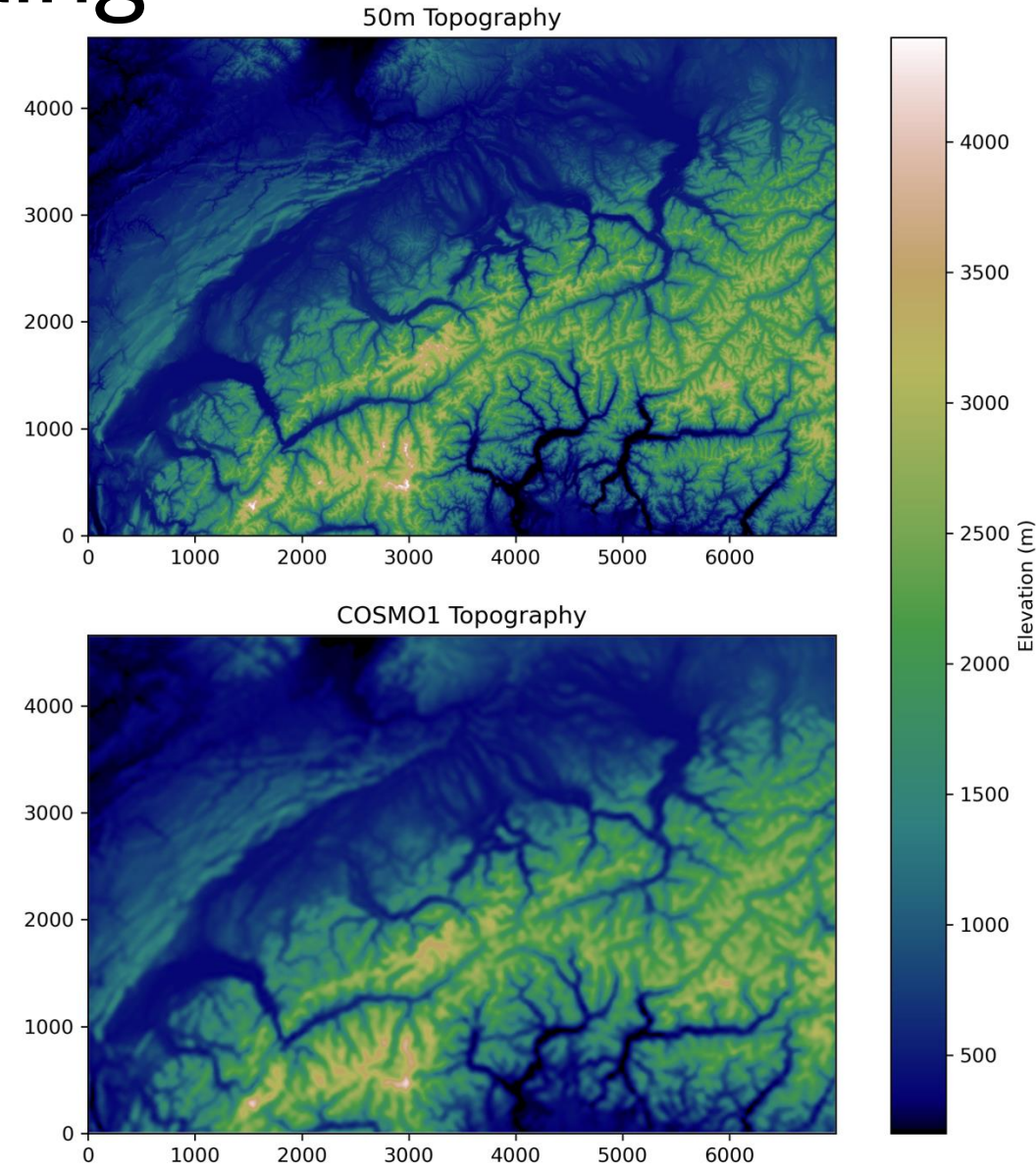


Zane Merva



Problem 2: Basic Downscaling

- Data from (old) Swiss weather model, COSMO1
 - Data for 11.03.2021 01:00
 - Model resolution of 1.1 km
- Digital Elevation Model (DEM) of Switzerland with 50m resolution
- Goal: High resolution maps of temperature and pressure



Basic Downscaling

