

Large-scale dynamics of the mid-latitude atmosphere : Part III

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Topics covered:

Two weeks ago

- Conservation of momentum: the equation of motion
- Hydrostatic balance, geostrophic balance
- Thermal wind balance

Last week

- Going beyond geostrophy: the ageotrophic wind
- Diagnosis of vertical motions with the ageostrophic wind: jet stream, trough/ridges
- Fronts and the Norwegian cyclone model

This week

· Extratropical cyclones: going beyond the polar front theory

Books



- J.E. Martin, «Mid-latitude atmospheric dynamics, a first course», 2006
- J.M. Wallace & P.V. Hobbs, «Atmospheric science, an introduction survey», 2006 →

JEM

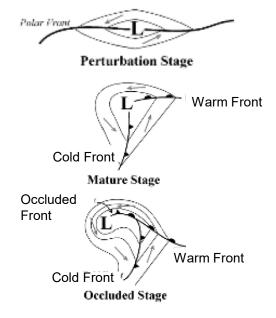
W&H

What are the limitations of the polar front theory?



Polar Front Cold, Polar Air
Warm, Tropical Air
Background State

- It assumes there is an existing front
- It does not explain why a perturbation along a frontal boundary will intensify
- It does not explain what leads to the dissipation of an extratropical cyclone (EC)
- It does not explain how the EC interacts with the upper-level dynamics
- → we are going to address some of these limitations today



Baroclinic instability and vertical structure of an extratropical cyclone



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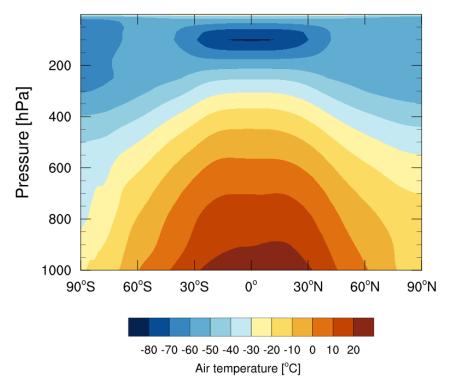
What does baroclinic mean?

- A baroclinic fluid is one in which temperature is not uniform on a constant pressure surface (i.e. density depends on both pressure and temperature).
- Differential solar heating between the poles and the equator creates a meridional temperature gradient
- → the temperature is not uniform on a constant pressure surface
- → the atmosphere is baroclinic

Quiz on seasonal and latitude dependence:

- Seasons: the atmosphere is most baroclinic in winter
- Geography: the atmosphere is most (least) baroclinic in the mid-latitudes (tropics)



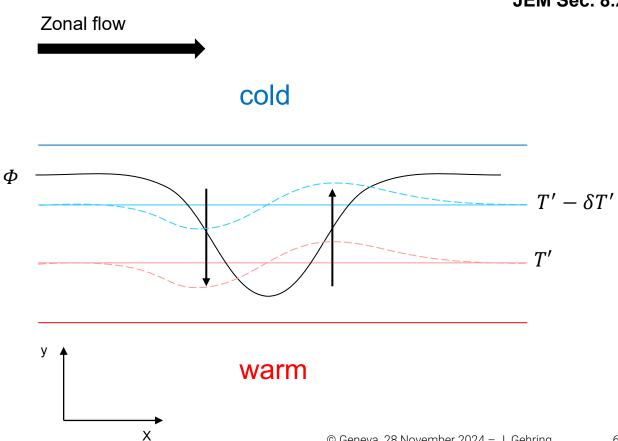


Why is a baroclinic atmosphere unstable?



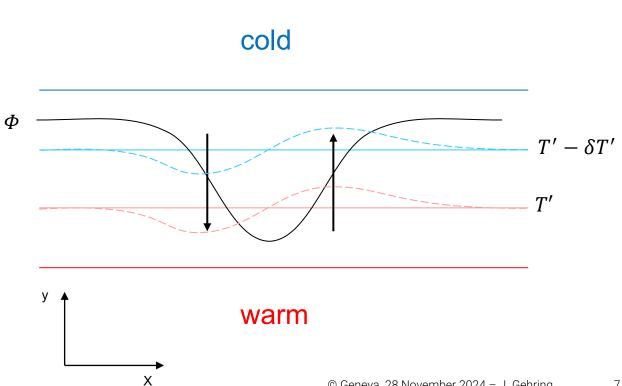
JEM Sec. 8.2

- Differential heating → meridional temperature gradient
- Wave-like perturbation in the geopotential height field
- Speed of this wave is the same as the background westerly flow
- Cold (warm) air advection upstream (downstream) of the trough axis



Is such a wave unstable?

- For the introduction of such a wave-like perturbation to be unstable:
 - The temperature anomalies must become larger
 - The kinetic energy of the wave motion must increase

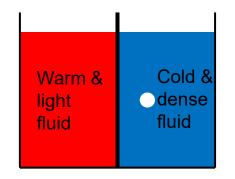


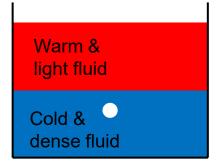
The energetic aspect



JEM Sec. 8.2

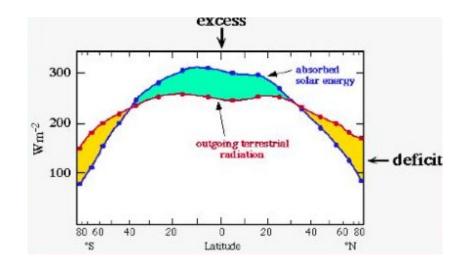
- The pole-to-equator temperature gradient represents a density contrast
- Centre of gravity at the mid-height point
- Remove the wall → potential energy converted into kinetic energy
- Lower center of gravity → lower potential energy → more stable





The role of extratropical cyclones in the atmospheric energy cycle

- Need a mechanism to redistribute heat meridionally
- Extratropical cyclones converts the available potential energy into kinetic energy
- Extratropical cyclones transport polar air equatorward and tropical air poleward -> redistribute the heat meridionally



How to get a surface low pressure?

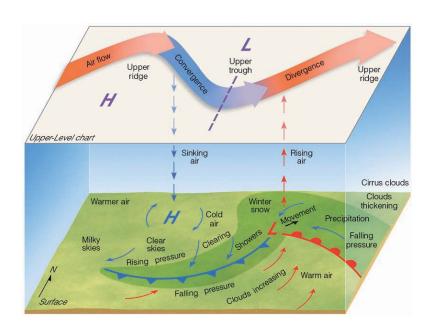


JEM Sec. 8.3

• Integrating the continuity equation $(\nabla_p \cdot \overrightarrow{V_h} = -\frac{\partial \omega}{\partial p})$ from p=0 to $p=p_s$ (the surface pressure) yields the **pressure tendency equation**:

$$\frac{\partial p_s}{\partial t} \approx -\int_0^{p_s} (\nabla \cdot \vec{V}) \partial p$$

- → The surface pressure tendency is the result of net mass convergence in the column
- → Net mass divergence (convergence) in the column leads to sea-level pressure fall (rise)
- Surface convergence in a low pressure system is the result of surface friction



From Ahrens, 2009, Meteorology Today

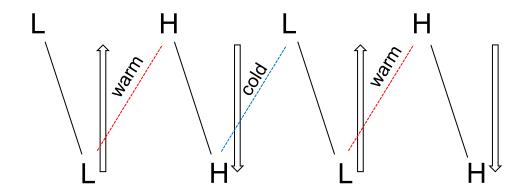
The vertical structure

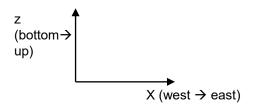


Pattern of **highs and lows**

JEM Sec. 8.2

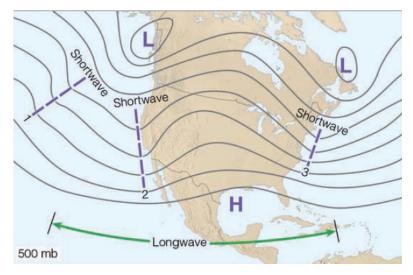
- Lows are associated with upper-level divergence and upward motions
- Upper-level divergence maximum is found ahead of the trough axis → westward tilt of the geopotential height axes
- thermal axes tilt eastward
- → Mid-latitude disturbances are characterised by thermally direct circulations in which warm air rises and cold air sinks





The most unstable wavelengths

- Spontaneous conversion of potential energy to kinetic energy → background zonal baroclinic flow is unstable to certain wave-like perturbation
- Wavelengths of 3000 to 4500 km are the most unstable.
- Corresponds to the wavelength of midlatitude shortwaves
- Shortwaves are very unpredictable, yet have a large influence on weather

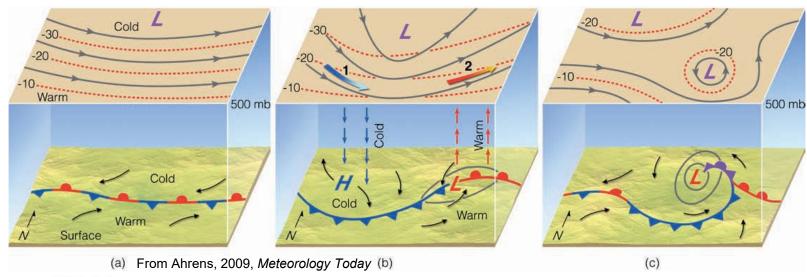


From Ahrens, 2009, Meteorology Today

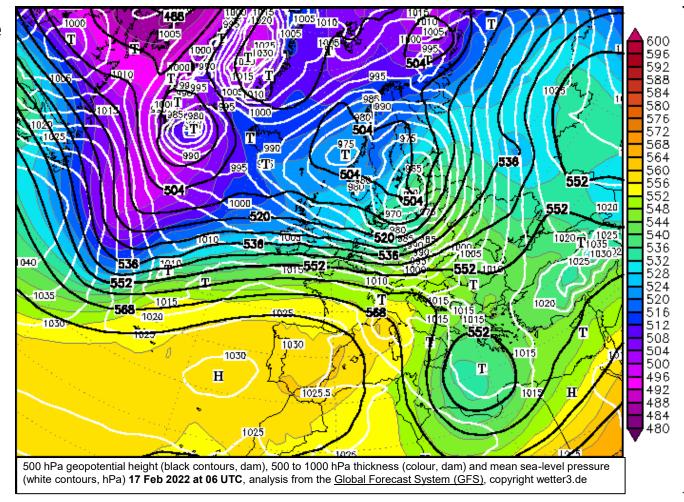
⁽¹⁾ Jules Charney (1947) and Eric Eady (1949) developed the theory of baroclinic instability independently and found the same result.

Cyclone decay

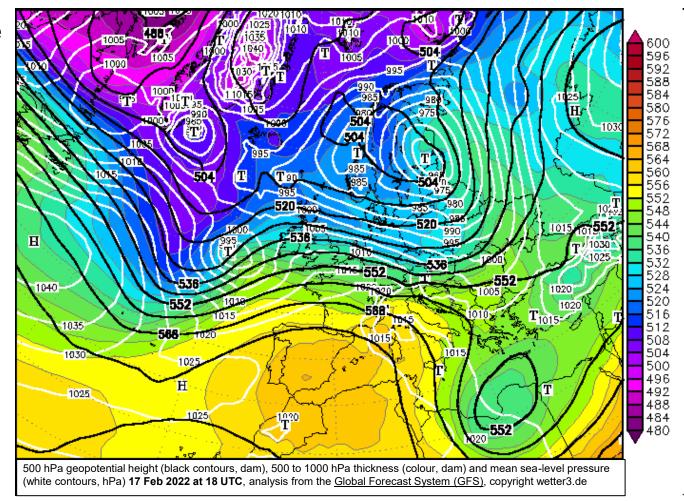
- Mature stage: westward tilt → maximum upper-level divergence above the sea level pressure minimum
- Dissipation stage: **geopotential height minima vertically aligned** → cyclone decays



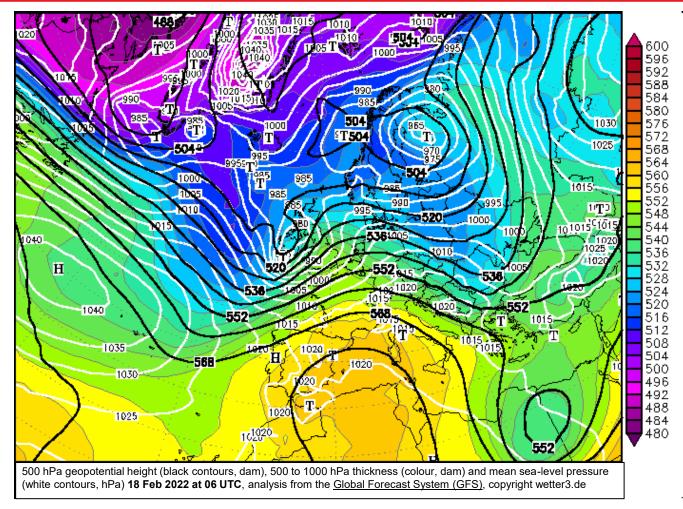
 The surface pressure trough is located ahead of the upper-level trough (westward tilt)



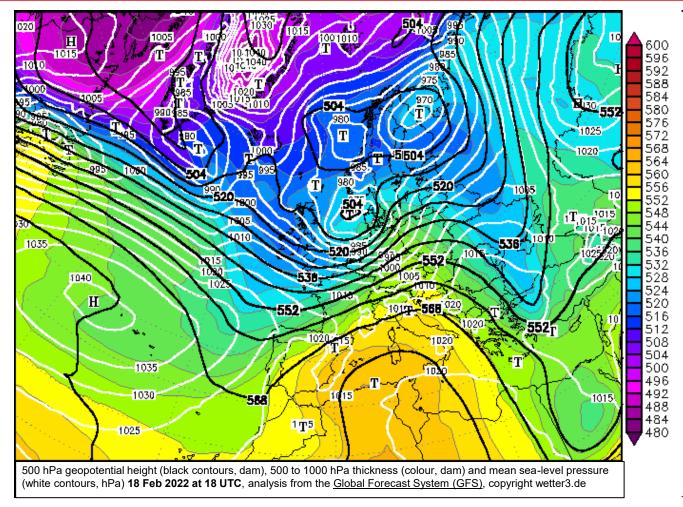
- The low pressure intensifies
- Westward tilt
- Warm air advection downstream of the trough axis strengthens the upper-level ridge



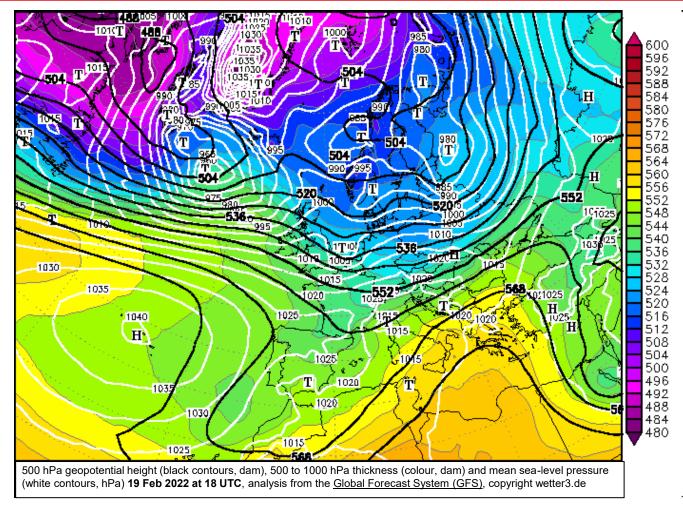
 Cyclone moves away from the peak of the warm sector: vertically aligned with the upper-level trough



 Cyclone perfectly aligned with the 500 hPa geopotential height minimum



Cyclone decays

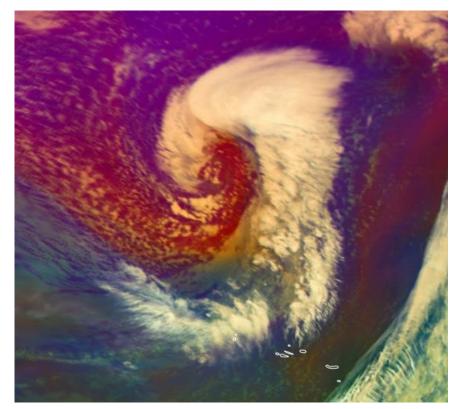


Baroclinic instability and vertical structure: take home messages

- Differential solar heating → meridional temperature gradient → baroclinic atmosphere
- Density contrast between the cold (dense) polar air and the warm (lighter) tropical air represents an unstable state
- A small perturbation of the background westerly flow will convert the available potential energy into kinetic energy
- Extratropical cyclones redistribute heat by transporting polar air equatorward and tropical air poleward
- Cyclone decay: it becomes vertically aligned with the upper-level trough → surface pressure rises

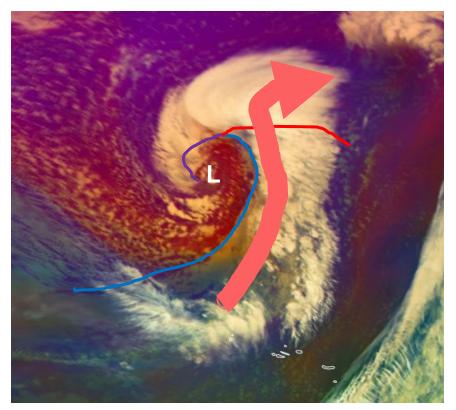
A Lagrangian perspective of extratropical cyclones: the conveyor belt model

- What explains the «comma» shape of the clouds around an extratropical cyclone?
- How does the air flows around the cyclone?



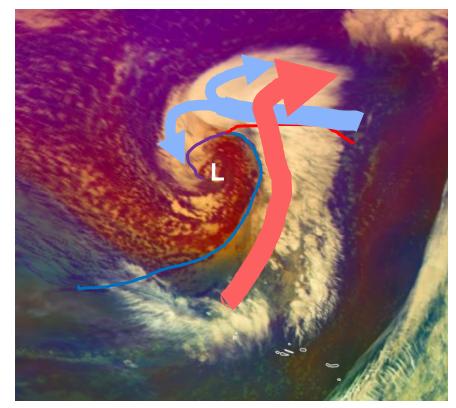
Satellite image of an extratropical cyclone over the northern Atlantic on 12 November at 00 UTC from Eumetsat airmass RGB product

- The warm conveyor belt (WCB): rising warm and moist air.
- WCBs are the primary cloud- and precipitationgenerating flow in extratropical cyclones and can lead to extreme precipitation
- Latent heat release intensifies the vertical motion and the horizontal temperature gradient (frontogenetic)



Satellite image of an extratropical cyclone over the northern Atlantic on 12 November at 00 UTC from Eumetsat airmass RGB product

 The cold conveyor belt: starts from the surface ahead of the warm front and rises slowly as it moves westward⁽¹⁾.



Satellite image of an extratropical cyclone over the northern Atlantic on 12 November at 00 UTC from Eumetsat airmass RGB product

(1) Relative to the cyclone's motion

• The dry conveyor belt:

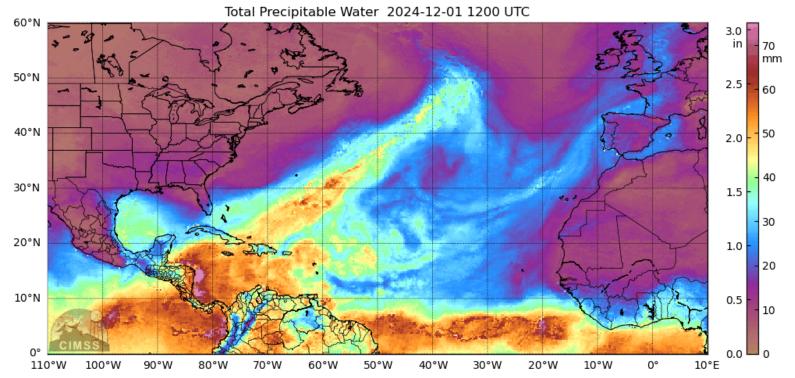
- Sinks behind the cold front → clear sky area in the cold sector.
- One branch turns towards the cyclone centre and leads to the dry slot, often visible in satellite images



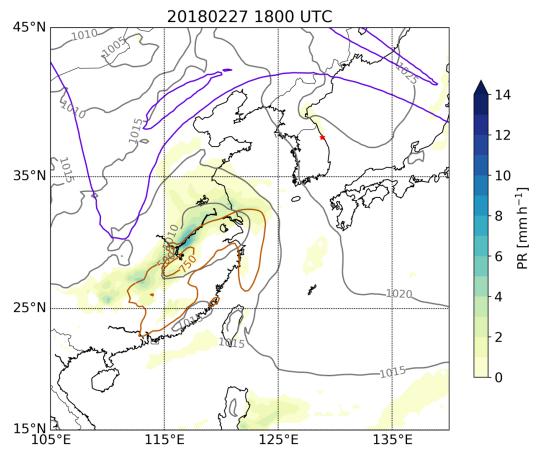
Satellite image of an extratropical cyclone over the northern Atlantic on 12 November at 00 UTC from Eumetsat airmass RGB product

The link between warm conveyor belts and atmospheric rivers

- Narrow filaments of enhanced water vapour transport: the **atmospheric rivers** (ARs)
- ARs are often associated with extratropical cyclones

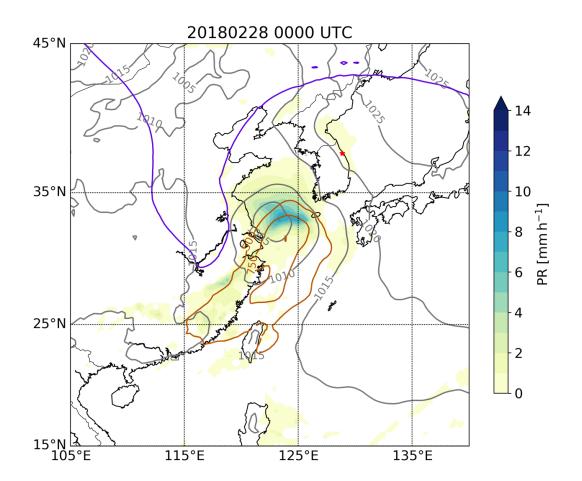


 An upper-level trough (blue contour) is located over eastern China with a surface low-pressure (grey contours) downstream

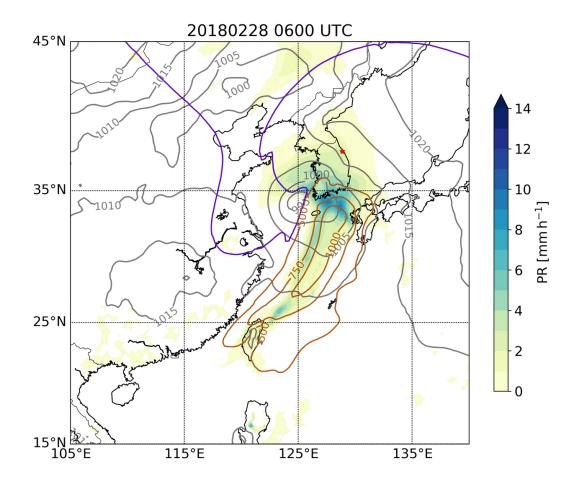


From Gehring et al. (2020)

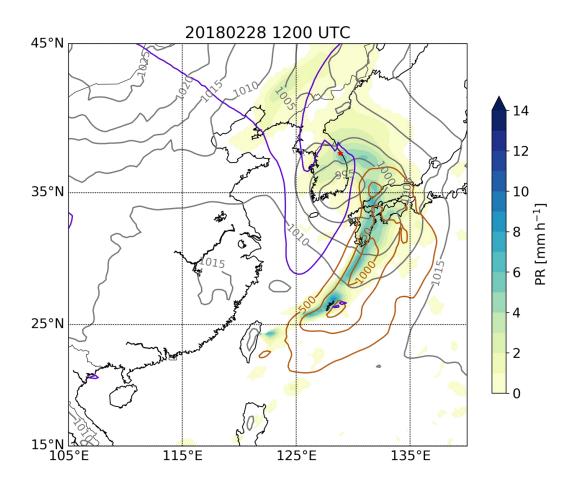
 The surface low intensifies as a result of the strong upper-level divergence



- The surface low further intensifies probably as a result of latent heat release in the warm conveyor belt (WCB)
- Intense moisture : the atmospheric river

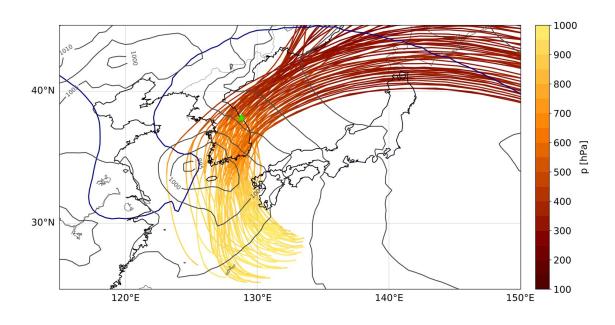


 The surface low becomes vertically aligned with the upper-level trough



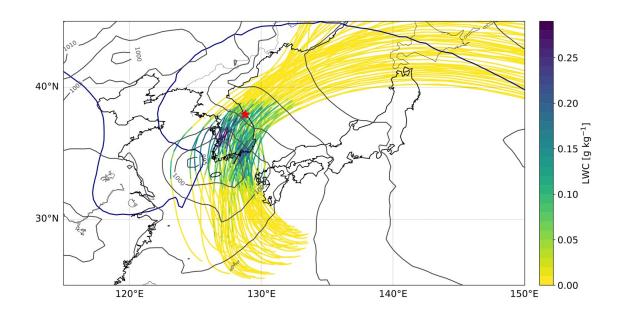
The warm conveyor belt shown with trajectory analysis

 Analysis of air trajectories shows a rising air mass ahead of the cold front



The warm conveyor belt shown with trajectory analysis

- The liquid water content (LWC) along the trajectories increases during the ascent as a result of condensation
- The latent heat release due to condensation intensifies the vertical motions and impact the upper-level dynamics



The conveyor belt model: take home messages



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- Analysis of air parcel trajectories in extratropical cyclones (ECs) reveal the presence of three
 distinguished airstreams with different temperature and humidity contents: the warm conveyor
 belt (WCB), the cold conveyor belt and the dry conveyor belt
- These conveyor belts explain the comma-shape structure of the clouds in an EC
- The airstream responsible for most of the cloud and precipitation in an EC is the ...
- Latent heat release in the WCB ... the upward motions and the horizontal temperature gradient
- WCBs and atmospheric rivers are often associated with each other and represent two complementary perspectives to study the moisture transport and precipitation associated with ECs